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Comparison of experimental and theoretical results for the structure and elastic properties of moganite

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Abstract: Moganite, which is monoclinic at ambient temperature, undergoes a displacive transition to an orthorhombic phase at ≈ 570 K. Whereas the monoclinic phase may be considered as α -quartz that is Brazil twinned along $\{1\ 0\ \bar{1}\ 1\}$ at the unit-cell scale (cell-twinning), the orthorhombic phase cannot be interpreted as a Brazil cell-twin of β -quartz, in contrast to statements made in the literature. The shape of the oxygen tetrahedra in monoclinic moganite has been determined more reliably by density functional theory (DFT) calculations than by experiment: the differences between the various experimental results for the shape of the oxygen tetrahedra at ambient temperature are typically ten times larger than the differences between the DFT results. The DFT calculations suggest that the oxygen tetrahedra in moganite are very close in shape to the oxygen tetrahedra in α -quartz. Among the three DFT calculations considered, the most convincing results for the bond angles in moganite are obtained for the DMol³ code with functional PBE.

Keywords: Brazil twinning; density functional theory; elastic properties; moganite structure.

Introduction

Moganite, a mineral detected as microcrystalline silica fillings of cavities and cracks in ignimbrite flows near the town of Mogán in the south of Gran Canaria, was first described by Flörke, Jones and Schmincke [1] as SiO₂-G. The name moganite was proposed by Flörke, Flörke and Giese [2]. After initial skepticism regarding the distinction between moganite and quartz, the “International

Commission on New Minerals and Mineral Names” approved moganite as a mineral species in 1999.

The structure of moganite was first determined by Miede et al. [3]. It has space group $C2/c$ (#15) and contains 12 Si atoms and 24 O atoms per conventional cell: Si1 at a position $4e$, Si2, O1, O2 and O3 at positions $8f$. Later structure determinations were made by Miede and Graetsch [4] and by Heaney and Post [5]. All these measurements were performed on powders of naturally occurring moganite, using X-ray diffraction and/or time-of-flight neutron diffraction. They confirmed the room temperature result given above but led to considerable differences in the structure parameters. An obvious reason for the differences is that no pure moganite has been found in nature.

Moganite may be considered as a Brazil cell-twin of α -quartz with composition plane $r = \{1\ 0\ \bar{1}\ 1\}$ and minimum lamellae thickness, i.e. thickness equal to the distance between neighboring $(1\ 0\ \bar{1}\ 1)$ planes [3]. Let \mathbf{a}_q , \mathbf{b}_q and \mathbf{c}_q define the usual primitive hexagonal cell of quartz. Then $\mathbf{a} = \mathbf{a}_q - \mathbf{b}_q$, $\mathbf{b} = \mathbf{a}_q + \mathbf{b}_q$, $\mathbf{c} = 2\mathbf{c}_q$, is a body-centered orthogonal cell for the Brazil cell-twin, for which its space group #15 appears in the setting $I2/a$. Note that the monoclinic angle β is equal to 90° for the unrelaxed cell-twin. To stress the structural relation between moganite and the quartz cell-twin, also the moganite structure is usually expressed in the setting $I2/a$, which has the advantage that the monoclinic angle β is close to 90° . Grimmer and Delley [6, 7] considered cell-twin models. For a given choice of the data for the quartz structure and given orientation of the composition plane, the models have a continuous degree of freedom, which corresponds to a translation between left- and right-handed quartz parallel to the monoclinic axis of the cell-twin. In particular, the translation may be chosen such that either $\angle(\text{O}-\text{Si}-\text{O})$ or $\angle(\text{Si}-\text{O}-\text{Si})$ has the same value as in quartz also across the composition plane. In the first case (model 1 of [6]) the oxygen tetrahedra are undistorted compared to quartz, in the second (model 3 of [6]) the angles between adjacent oxygen tetrahedra are as in quartz. It turns out that model 1 is closer to the experimental results.

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Calculations of the moganite structure using density functional theory (DFT) were done by Hantsch et al. [8] and by us for the present paper. They correspond to moganite at a temperature of 0 K and will be compared to the experimental data at ambient temperature and to cell-twin model 1, based either on the quartz data of Lager et al. [9] at 13 K or on those of Baur [10] at 291 K.

Moganite, its structure determined by theory and experiment

Brazil cell-twin models

Figure 1 illustrates the structure of moganite by showing the cell-twin model based on quartz data [10] at 291 K projected parallel to its monoclinic axis **b**. O1 binds to two Si2, O2 (and O3) to a Si1 and a Si2. O3 connects two oxygen tetrahedra, the central Si atoms of which lie on opposite sides of a Brazil twin boundary (marked in brown). The bonds are indicated by arrows. The atom at the arrow-head has the larger *y* coordinate. The figure shows the situation for model 1; the figure for model 3 is identical except that the arrows that cross a Brazil twin boundary have their head at the opposite end of the dark green line. The corresponding figure for the cell-twin based on the quartz data of Lager et al. [9] at 13 K looks practically the same; the corresponding figures for the experimental and DFT results mentioned in the introduction differ only slightly; for all of them the sense of the arrows is the same as for cell-twin model 3. This tells us that, as far as possible, relaxation conserves the angles (O–Si–O) as well as (Si–O–Si) of quartz.

Columns M1,2 of Table 1 describe unrelaxed Brazil cell-twins, whose tetrahedra formed by 4 O atoms bound to the same Si atom have the same bond distances and bond angles as the corresponding tetrahedra in α -quartz. Only the Si–O–Si angles between two tetrahedra having O in common and the two Si on opposite sides of a twin boundary have values different from the values for α -quartz. The model given in column M2 corresponds to model 1 in Table 2 of [6]; Grimmer and Delley [7] show how the value $Y=2\delta=0.4395$ given for this model is related to the fault vector **f** of Lang [11] and the displacement vector **R** of Phakey [12], in particular $|\mathbf{f}|=0.4395 a$, where *a*, *c* are the lattice parameters of quartz. Column M1 is based on low temperature data for quartz, which lead to a fault vector $|\mathbf{f}|=0.4464 a$.

Figure 2 gives a graphical representation of the results in the lower part of Table 1.

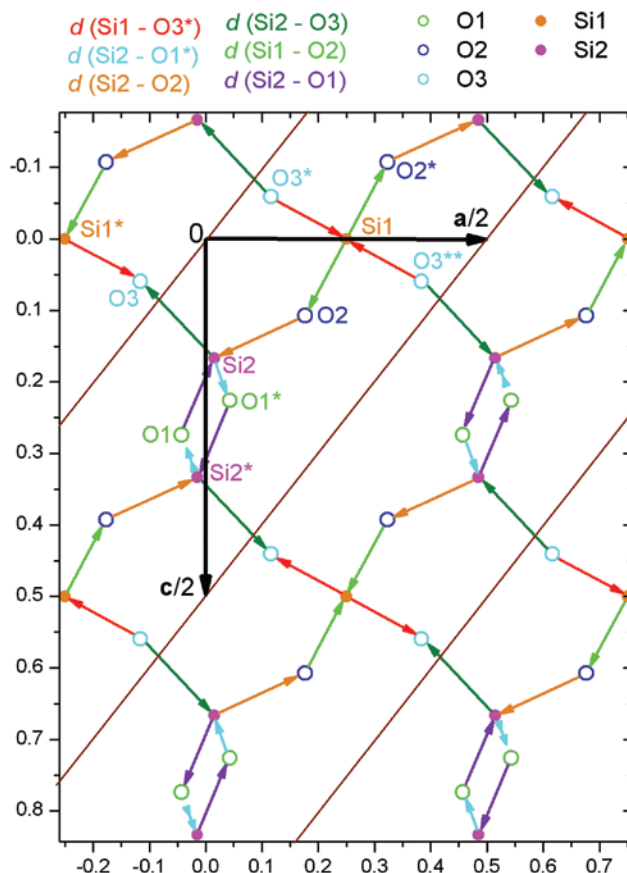


Fig. 1: A conventional *I*-centered monoclinic cell of moganite projected parallel to its monoclinic axis **b**. The named atoms are those at positions $xa + yb + zc$, whose coordinates (*x*, *y*, *z*) are obtained from column M2 of Table 1 as follows: Si1($\frac{1}{4}$, *y*1, 0), Si1*($-\frac{1}{4}$, $-y$ 1, 0), Si2(*x*2, *y*2, *z*2), Si2*($-x$ 2, y 2 + $\frac{1}{2}$, $-z$ 2 + $\frac{1}{2}$), O1(*X*1, *Y*1, *Z*1), O1*($-X$ 1, Y 1 + $\frac{1}{2}$, $-Z$ 1 + $\frac{1}{2}$), O2(*X*2, *Y*2, *Z*2), O2*($-X$ 2 + $\frac{1}{2}$, *Y*2, $-Z$ 2), O3(*X*3, *Y*3, *Z*3), O3*($-X$ 3, $-Y$ 3, $-Z$ 3), O3**(X 3 + $\frac{1}{2}$, $-Y$ 3, *Z*3). Si and O atoms connected by arrows of the same color have the same distance; this holds not only for M2 but for all eight cases considered in Table 1. The atom at the arrow-head has the larger *y* coordinate; double arrows start at an O-atom in the unit cell below or end at an O-atom in the unit cell above. The brown lines show the twin boundaries of the Brazil cell-twin models considered in Table 1. The sense of the dark green arrows holds for both cell-twin models, M1 and M2; it is opposite for all the experimental and DFT results given in the last six columns of Table 1.

Experimental results

The structure of moganite has been determined experimentally in [3–5]. So far, it has not been possible to synthesize moganite as a pure phase; in nature it is found intergrown with fine-grained quartz and containing small amounts of volatile (H_2O , CO_2) and non-volatile (Na_2O , K_2O , Fe_2O_3 , ...) impurities. Relatively pure moganite from Gran Canaria was used in [3–5]. Typical impurities are

Tab. 1: Cell parameters a , b , c , β and position parameters of moganite in $1/2/a$ 1 setting; cell volume V , bond distances d and bond angles according to various authors. In columns M1,2 the Brazil cell-twin model 1 of Grimmer and Delley [6, 7] with minimum lamellae thickness ($N=1$) is interpreted as moganite. Column E2 uses for O2 the position parameters adjusted by distance least squares, not the ones obtained by unconstrained structure refinement.

	Brazil twin models		Experimental results			DFT results		
	M1 using 13 K data [9]	M2 using 291 K data [10]	E1 Miehe et al. [3]	E2 Miehe and Graetsch [4]	E3 Heaney and Post [5]	D1 Hantsch et al. [8]	D2 DMol ³ PBE	D3 DMol ³ PBEsol
a [Å]	4.9021·√3	4.913·√3	8.770(2)	8.758(2)	8.7371(6)	8.903	8.792	8.611
b [Å]	4.9021	4.913	4.879(1)	4.876(1)	4.8692(3)	4.995	4.976	4.861
c [Å]	10.7994	10.809	10.720(2)	10.715(2)	10.7217(7)	10.907	10.877	10.735
β [°]	90	90	90.08(4)	90.08(3)	90.193(9)	90.36	90.64	90.52
V [Å ³]	449.48	451.9	458.7(3)	457.6(3)	456.12(5)	485.0	475.8	449.3
Si1 in 4(e)								
x_1	1/4	1/4	1/4	1/4	1/4	1/4	1/4	1/4
y_1	-0.0588	-0.06035	-0.0228(20)	-0.0092(17)	-0.0274(11)	-0.0177	-0.0315	-0.0382
z_1	0	0	0	0	0	0	0	0
Si2 in 8(f)								
x_2	0.0160	0.01505	0.0072(5)	0.0115(4)	0.0103(3)	0.0077	0.0144	0.0174
y_2	0.2392	0.23480	0.2507(7)	0.2533(6)	0.2486(7)	0.2363	0.2399	0.2527
z_2	1/6	1/6	0.1688(4)	0.1678(2)	0.1682(4)	0.1668	0.1669	0.1666
O1 in 8(f)								
X_1	-0.0438	-0.04305	-0.0196(9)	-0.0314(12)	-0.0219(3)	-0.0142	-0.0268	-0.0326
Y_1	0.0382	0.03010	0.0701(13)	0.0680(13)	0.0644(5)	0.0406	0.0505	0.0712
Z_1	0.275183	0.273933	0.2923(7)	0.2860(5)	0.2878(2)	0.2840	0.2852	0.2879
O2 in 8(f)								
X_2	0.1794	0.17670	0.1667(7)	0.1711	0.1678(3)	0.1627	0.1785	0.1847
Y_2	0.1314	0.12915	0.1671(12)	0.1770	0.1708(5)	0.1703	0.1582	0.1554
Z_2	0.108517	0.107267	0.1030(6)	0.1050	0.1002(3)	0.0969	0.1081	0.1123
O3 in 8(f)								
X_3	-0.1144	-0.11635	-0.1274(6)	-0.1343(5)	-0.1297(3)	-0.1321	-0.1184	-0.1131
Y_3	1/4	1/4	0.2217(19)	0.2148(12)	0.2296(5)	0.2062	0.2236	0.2342
Z_3	0.05815	0.05940	0.0668(6)	0.0739(8)	0.0675(3)	0.0727	0.0605	0.0561
1 $d(\text{Si1-O2})$ [Å]	1.6131	1.6125	1.6166	1.6030	1.6145	1.6168	1.6383	1.6331
2 $d(\text{Si1-O3}^*)$ [Å]	1.6120	1.6043	1.6151	1.6299	1.6102	1.6144	1.6336	1.6266
3 $d(\text{Si2-O1}^*)$ [Å]	1.6120	1.6043	1.6168	1.6217	1.6115	1.6128	1.6343	1.6282
4 $d(\text{Si2-O2})$ [Å]	1.6120	1.6043	1.6197	1.5961	1.6048	1.6147	1.6358	1.6293
5 $d(\text{Si2-O1})$ [Å]	1.6131	1.6125	1.6079	1.6009	1.5907	1.6214	1.6390	1.6339
6 $d(\text{Si2-O3})$ [Å]	1.6131	1.6125	1.6141	1.6354	1.6314	1.6155	1.6365	1.6296
7 $\angle(\text{O2-Si1-O3}^*)$, $\angle(\text{O2}^*-\text{Si1-O3}^{**})$ [°]	110.7	110.4	110.2	114.9	111.9	110.3	110.5	110.9
8 $\angle(\text{O2}^*-\text{Si1-O2})$ [°]	109.4	109.5	110.1	111.0	106.6	109.0	109.6	109.6
9 $\angle(\text{O2-Si1-O3}^{**})$, $\angle(\text{O2}^*-\text{Si1-O3}^*)$ [°]	108.6	108.8	110.0	106.0	111.0	109.3	108.9	108.5
10 $\angle(\text{O3}^*-\text{Si1-O3}^{**})$ [°]	108.9	109.0	106.1	104.1	104.6	108.6	108.4	108.3
11 $\angle(\text{O1}^*-\text{Si2-O1})$ [°]	108.6	108.8	109.0	108.5	108.2	108.1	107.9	107.2
12 $\angle(\text{O1}^*-\text{Si2-O2})$ [°]	108.9	109.0	107.3	104.8	107.8	108.7	107.7	108.3
13 $\angle(\text{O1-Si2-O2})$ [°]	110.7	110.4	110.4	114.1	112.8	110.9	111.6	111.7
14 $\angle(\text{O1-Si2-O3})$ [°]	109.4	109.5	113.7	103.8	111.5	110.4	111.4	111.4
15 $\angle(\text{O2-Si2-O3})$ [°]	108.6	108.8	108.3	113.4	109.2	109.8	109.7	109.2
16 $\angle(\text{O1}^*-\text{Si2-O3})$ [°]	108.9	110.4	107.9	112.3	107.1	109.0	108.5	108.9
17 $\angle(\text{Si2-O1-Si2}^*)$ [°]	142.4	143.7	136.9	138.6	139.7	145.7	141.2	137.0
18 $\angle(\text{Si1-O2-Si2})$ [°]	142.4	143.7	146.4	143.8	145.8	147.5	140.4	138.0
19 $\angle(\text{Si1}^*-\text{O3-Si2})$ [°]	139.0	138.8	147.0	148.6	145.0	149.6	144.2	142.6

1–3 % H₂O, 0.2–1 % CO₂, and less than 0.2 % non-volatile impurities by weight [4]. Powder diffraction and Rietveld refinement were used in [3–5] for structure determination.

Miehe et al. [3] used time-of-flight (TOF) neutron scattering and X-ray diffraction, Miehe and Graetsch [4] used Cu K α X-ray diffraction, Heaney and Post [5] used TOF

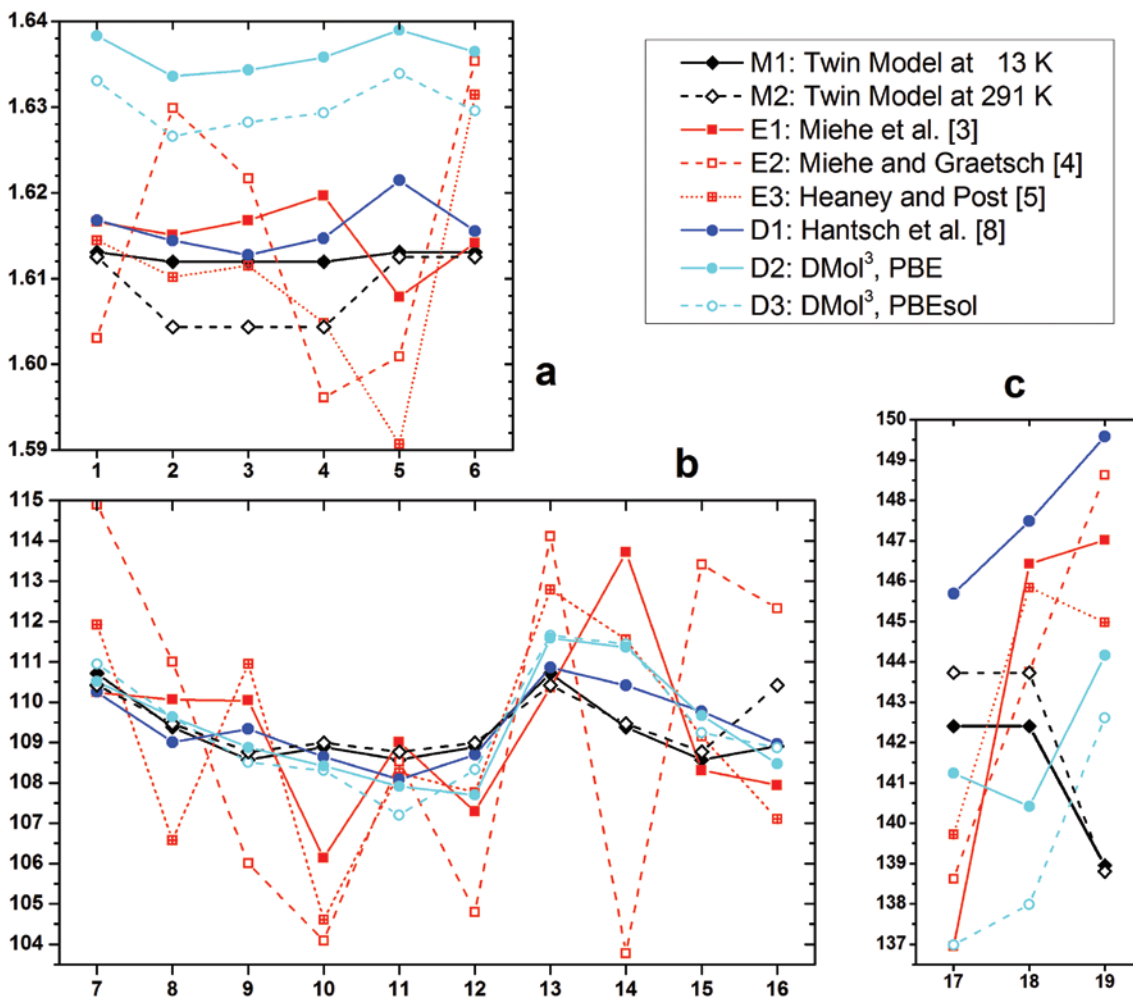


Fig. 2: (a) Shows Si–O distances in Å, (b) O–Si–O angles and (c) Si–O–Si angles in °. The lines are a guide to the eye. The numbers 1–19 refer to the corresponding rows of Table 1. Experimental values are shown in red, DFT results in blue or cyan. The values in black correspond to Brazil cell-twin model 1 with minimum lamellae thickness of Grimmer and Delley [6, 7] based on quartz data at 13 K [9] and at 291 K [10], respectively. Distance 6 as well as angles 14–16 and 19 correspond to atoms on different sides of Brazil twin boundaries. Distances 1–2 and angles 7–10 involve Si1, distances 3–6 and angles 11–16 involve Si2.

neutron scattering at 298 K. Their results are given in columns E1–E3 of Table 1.

Cell parameters and cell volume of moganite can be determined more reliably by experiment than by Brazil twin models or DFT calculations. Nevertheless, the differences between the values given in columns E1–E3 of Table 1 are in most cases much larger than the experimental uncertainties given by the authors, as shown in Figure 3. According to [5], the actual errors may be more than an order of magnitude higher than the deviations shown, which were obtained by the General Structure Analysis System (GSAS) [13].

Let us denote the values in the various columns of Table 1 by the corresponding subscript. Cell parameters and cell volume can be determined reliably by experiment. Neglecting deviations from the nominal $\text{Si}_{12}\text{O}_{24}$ cell

content, the density ρ will be inversely proportional to the cell volume V . The density of α -quartz at room temperature corresponds to the cell volume V_{M2} . The experimentally determined cell volumes of moganite, V_{E1} , V_{E2} and V_{E3} being larger than V_{M2} , moganite will have slightly smaller density than quartz.

Also for the position parameters, the differences between the values given in columns E1–E3 of Table 1 are in most cases much larger than the experimental uncertainties given by the authors, as shown in Figure 4.

Whereas unconstrained structure refinement led to plausible Si–O distances for the neutron data [5], it produced for the Cu $K\alpha$ data [4] an O2 position unrealistically far from Si1 (1.643 Å) and unrealistically close to Si2 (1.537 Å). The E2 position parameters for O2 were therefore adjusted by minimizing the sum of the squared

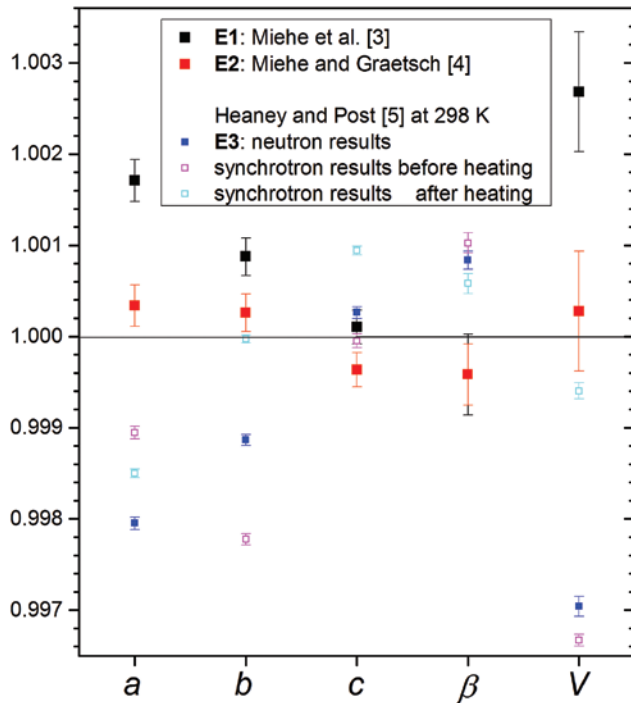


Fig. 3: Experimental values of the cell parameters and cell volume of moganite given in columns E1, E2 and E3 of Table 1 divided by their average. Synchrotron powder X-ray diffraction results at 298 K before and after heating to 1354 K are also given [5]. These values were not considered for computing the average.

Si–O2 distances. The results of Heaney and Post [5] can be expected to be the most reliable experimental results because structure refinement based on their powder neutron diffraction data was possible without restraints.

Figures 2a and b show that the shapes of the O-tetrahedra are difficult to determine experimentally. The results of [3] and [4] differ widely: whereas according to [3] $d(\text{Si}2\text{--O}2)$ (#4) is the largest among the Si–O distances and $\angle(\text{O}1\text{--Si}2\text{--O}3)$ (#14) the largest among the O–Si–O angles, they are the smallest distance and smallest angle according to [4], the results of [5] lying in between. These discrepancies suggest that the true Si–O distances #1–5, the true O–Si–O angles #7–13 and the true Si–O–Si angles #17–18 will be closer to the cell-twin model values than to the experimental ones. Relaxation of the atom positions in the models will mainly affect the remaining distances and angles, which involve atoms on both sides of a Brazil twin boundary.

Density functional theory calculations

Columns D2 and D3 of Table 1 give the results of density functional theory (DFT) calculations with the DMol³ code

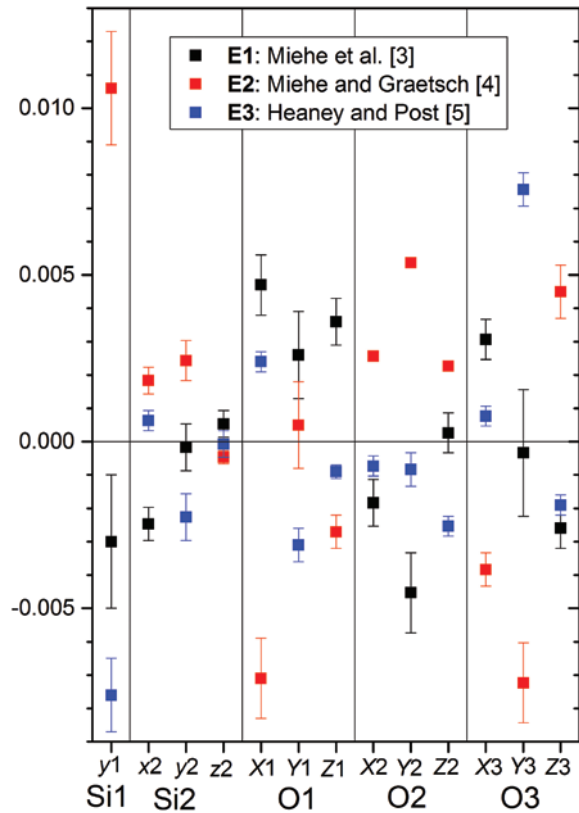


Fig. 4: Difference between the experimental values of the position parameters given in columns E1, E2 and E3 of Table 1 and the average of these 3 values. (Note that no uncertainties are given for O2 in E2).

[14, 15], using the functional PBE as defined in [16] for D2 and the functional PBEsol as defined in [17] for D3. Whereas PBEsol tends to underestimate the cell volume (V_{D3} is 1.8 % smaller than V_{E2}), PBE tends to overestimate it (V_{D2} is 4.0 % larger than V_{E2}). In both cases a $4 \times 4 \times 4$ Γ -centered mesh in \mathbf{k} -space was used as well as the default local orbital basis set DNP [15] with cutoff radii for the basic functions equal to $7.32 a_0$ for O and $10.06 a_0$ for Si, a_0 denoting the Bohr radius. The calculations minimized energy with respect to the parameters defining the moganite cell and the Wyckoff positions, starting out with the moganite structure given in [4]. The results of Hantsch et al. [8] are given for comparison in column D1. Also they used the functional PBE.

As discussed in Section Experimental results, we can assume that the true Si–O distances in moganite at room temperature are close to the experimental values for α -quartz, which appear in the Brazil cell-twin model M2 of moganite. Although the parameters for the moganite cell obtained in D1 are on average 2% larger than the experimental values, the Si–O distances are only 0.5% larger than the Si–O distances in quartz. This means that

the distances between neighboring oxygen tetrahedra and, therefore, the Si–O–Si angles must be larger than in reality. The opposite situation occurs for D3, where the parameters for the moganite cell are on average 0.6% too small and the Si–O distances 1.4% larger than for quartz. This means that the Si–O–Si angles must be smaller than in reality. For D2 the parameters for the moganite cell are on average 1.3% too large and the Si–O distances 1.7% larger than for quartz, so that the Si–O–Si angles will be only little smaller than in reality. This situation is clearly shown in Figure 2c. Table 1 shows that all the 13 independent Wyckoff parameters D2 lie between the corresponding values D1 and D3, the only exception being z_2 , for which all three values are almost equal.

Also the other suggestion made in Section Experimental results that the O–Si–O and Si–O–Si angles obtained in the Brazil cell-twin models are close to reality for the cases where the atoms involved are not separated by a Brazil twin boundary is substantiated by the DFT results: The DFT results D1–D3 for the O–Si–O angles #7–13 and the D2 results for the Si–O–Si angles #17–18 are close to the cell-twin model results. For the O–Si–O angle #14 and, in particular, the Si–O–Si angle #19 the DFT results D1–D3 deviate more strongly from the results M1–M2. Figure 2b clearly shows that the results of the three DFT calculations give more consistent results than the experiments E1–E3. We conclude that the DFT results, and in particular D2 give a better description of the Wyckoff position parameters and O–Si–O and Si–O–Si angles of moganite than the experimental results E1–E3.

As for the experimental results shown in Figure 3, the differences of the various DFT results are largest for lattice parameter a . The situation is similar for the position parameters of Si: the DFT results differ most for the y parameter of Si1 and least for the z parameter of Si2, as is the case for the experimental results shown in Figure 4.

Comparing the position parameters of Si1, Si2, O1, O2 and O3 in column D3 of Table 1 with the corresponding parameters in either column M1 or M2 we see that our DFT results differ from the two Brazil cell-twin models mainly in the y -coordinate, i.e. perpendicular to the plane shown in Figure 1. The difference in the y -coordinate has opposite sign for O3 than for Si1, Si2, O1 and O2. The largest shift in each of x , y and z has O1.

Denote by ΔE the bond enthalpy per mole SiO_2 of moganite compared to α -quartz. Hantsch et al. [8] obtained $\Delta E_{D1} = -0.35$ kJ/mol, i.e. moganite should be the stable phase according to the sign of ΔE_{D1} . However, its absolute value is less than the uncertainty of their calculations, which they estimate at 4 kJ/mol. Taking lattice and position parameter relaxation into account, we

obtained $\Delta E_{D2} = 0.4$ kJ/mol and $\Delta E_{D3} = 0.7$ kJ/mol, stating that α -quartz is the stable phase at low temperature and pressure. These values are smaller than the experimental CODATA value at 25°C and atmospheric pressure $\Delta E = 3.4 \pm 0.7$ kJ/mol [18].

α - and β -moganite

Heaney and Post [5] also investigated the dependence of the lattice parameters on temperature between 100 and 1354 K using synchrotron X-ray diffraction. They found a reversible displacive phase transition from monoclinic α -moganite to orthorhombic β -moganite at 569 K. Whereas in [5] it is argued that the transition from α - to β -moganite is second-order, hard-mode Raman spectroscopy results indicate that the situation is more involved [19].

Similarly as in [4], constraints were used also in [5] for structure refinement from their X-ray data: all Si–O bonds were assumed to be of length 1.61 Å in β -moganite. Table 2 shows that the structure given in Tables 4 and 5 of [5] for β -moganite at 1354 K is close to the structure proposed in [20] for moganite at room temperature before it was recognized that moganite has only monoclinic symmetry at ambient conditions [3]. Following Heaney and Post [5] we use the setting $Imab$ of space group $Ibam$ (#72) to facilitate comparison with monoclinic moganite. The atom position parameters given in [5] for α -moganite at 298 K and for β -moganite at 1354 K differ at most by 0.03.

Figure 5 illustrates the structure of α -moganite as obtained by the DMol³ code with functional PBE and β -moganite as determined in [5].

It is interesting to compare the transition in moganite with the transition from α - to β -quartz, which takes place

Tab. 2: The structure of orthorhombic moganite (with space group $Ibam$) in $Imab$ setting.

Cell and position parameters	Heaney and Post [5] for moganite at 1354 K	Miehe et. al. [20] for moganite at room temperature
a [Å]	8.8159(9)	8.74
b [Å]	4.9371(5)	4.88
c [Å]	10.7605(14)	10.70
Si1 in $4a$	$\frac{1}{4}, 0, 0$	$\frac{1}{4}, 0, 0$
Si2 in $8j$	0, 0.2255(6), 0.1701(3)	0, 0.250, 0.170
O1 in $8j$	0, 0.0416(7), 0.2935(4)	0, 0.065, 0.290
O2 in $16k$	0.1471(3), 0.2007(6), 0.0824(5)	0.147, 0.189, 0.086
$d(\text{Si-O})$ [Å]	1.61	1.57–1.60
$\angle(\text{O-Si-O})$ [°]	102.5–116.2	105.4–110.7
$\angle(\text{Si-O-Si})$ [°]	138.5, 145.5	140.7, 153.3

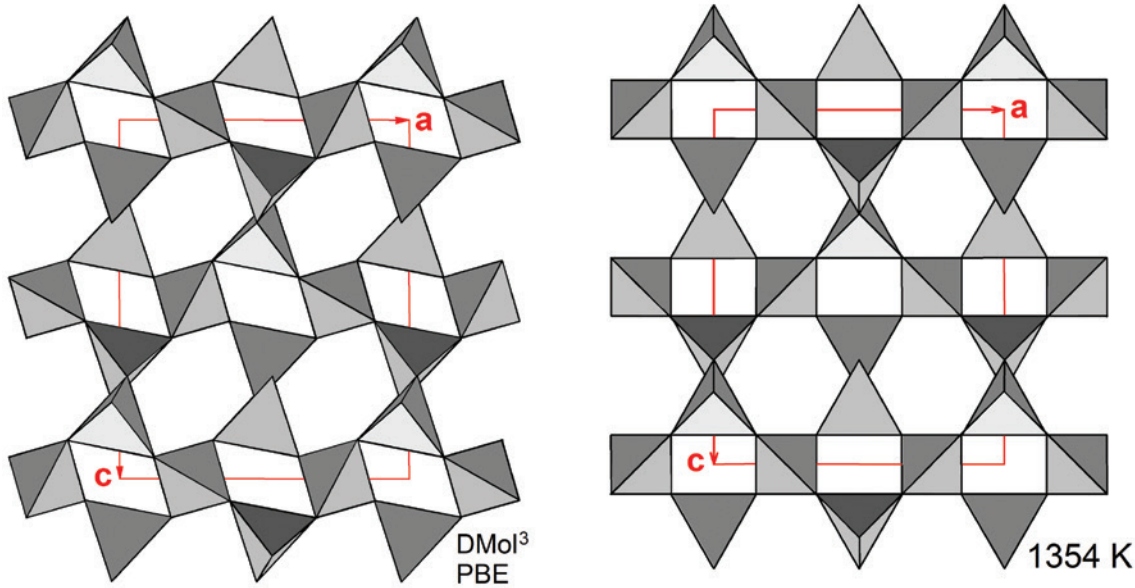


Fig. 5: Oxygen tetrahedra of α -moganite (left) with space group #15 in $I2/a$ setting and β -moganite (right) with space group #72 in $Imab$ setting.

at 847 K and is of first order [21]. Whereas the structure of α -moganite can be interpreted as Brazil cell-twinned α -quartz, β -moganite cannot be interpreted as Brazil cell-twinned β -quartz although the contrary is stated in [20] and suggested in Table 5 of [4]. In fact, a periodically Brazil twinned α - or β -quartz cannot have space group $Imam$; stacking $(1\ 0\ \bar{1}\ 1)$ layers alternately of left- and right-handed quartz mentioned in [20] produces monoclinic, not orthorhombic moganite.

Using molecular dynamics simulations Murashov and Svishchev [22] found that at ambient temperature and pressure β -moganite has the lowest enthalpy among pure silica phases, 5.5 kJ/mol lower than α -quartz. This is in contradiction with experiment and DFT results. (Note that it was for α -moganite that a slightly lower enthalpy (0.35 kJ/mol) than for α -quartz was obtained in [8].) Also the distribution of Si–O–Si angles in β -moganite given in Figure 3 of [22] shows unrealistic maxima at 130° and 159° , which contrast with the values 138.5° and 145.5° obtained in [5]. No phase transition of moganite at ambient pressure was found in [22] for the temperature range from 100 K to 1100 K in contrast to experiment [5].

Elastic properties

In their investigation of the high-pressure behavior of moganite by angle-dispersive X-ray powder diffraction, Léger et al. [23] determined the bulk modulus from the

pressure dependence of the molar volume, obtaining $\kappa = 32.2$ GPa.

Hantsch et al. [8] compared the stiffness constants obtained by their DFT calculations with the results from lattice energy minimization calculations using the program GULP [24] either with rigid ion BKS potentials [25] or with shell model SG potentials [26]. Their results for moganite are presented in Table 3 together with our results. The DMol³ code [14, 15] was used with basis set DNP [15] and either functional PBE [16] or PBEsol [17] to

Tab. 3: Elastic stiffness constants c_{ij} and bulk modulus κ of moganite according to various authors.

[GPa]	Hantsch et al. [8]			This paper	
	GULP BKS	GULP SG	D1 DFT	D2 DMol ³ PBE	D3 DMol ³ PBEsol
c_{11}	142.5	84	88	71.11	67.18
c_{22}	111.3	60	45	55.81	56.52
c_{33}	157.8	74	117	87.41	97.21
c_{44}	34.9	14	15	31.21	32.90
c_{55}	46.2	43	40	46.37	40.35
c_{66}	40.3	22	6	14.56	10.86
c_{12}	24.2	5	−7	−7.26	−7.45
c_{13}	59.1	3	27	7.57	4.46
c_{15}	−18.5	4	−8	−1.50	3.31
c_{23}	24.3	−17	−12	3.58	6.31
c_{25}	0.2	7	−9	7.09	8.35
c_{35}	−18.9	3	−5	−7.13	−5.46
c_{46}	13.2	4	1	−2.72	−3.47
κ	66.7	21	20	22.9	22.5

determine the stress in strained moganite. The stiffness constants in the last two columns of Table 3 were obtained by numerical differentiation of the stresses as functions of the strains.

The results D2 and D3 obtained with the DMol³ code but different functionals agree quite well. In the discussion of their results Hantsch et al. [8] argue that their GULP BKS results for moganite are less reliable than their GULP SG and DFT results and that the bulk modulus κ is smaller for moganite than for α -quartz (for which $\kappa \approx 37.4$ GPa [27, 28]) but higher than their GULP SG and DFT values. Our results confirm these conclusions but our bulk modulus is still considerably smaller than the value $\kappa = 32.2$ GPa of Léger et al. [23].

Conclusion

The oxygen tetrahedra in our Brazil cell-twin models M1 and M2 (see Table 1) coincide in shape and size with those of quartz, i.e. their Si–O distances and O–Si–O angles agree with those of quartz at 13 K and 291 K, respectively. Although the experimentally determined lattice parameters a , b , c and β differ for E1, E2 and E3 much more than the given uncertainties, these parameters can be determined experimentally much more precisely than by means of DFT calculations, the results of which strongly depend on the functional employed. The situation is completely different for the Si–O distances and O–Si–O angles, which determine the shape and size of the oxygen tetrahedra. The DFT calculations D1, D2 and D3 agree that the Si–O distances decrease in the order #5, #1, #6 and #4; #2,3 being the smallest distances. For each of the ten O–Si–O angles (#7–16), the DFT calculations D1, D2 and D3 give similar results, which are close to those for the Brazil cell-twin models M1 and M2, especially in the cases, where O–Si–O does not cross a Brazil twin boundary. If O–Si–O does cross a Brazil twin boundary, one expects atomic relaxation to modify the O–Si–O angles more strongly, as confirmed by the DFT results. The experimentally determined Si–O distances and O–Si–O angles differ wildly for E1, E2 and E3, showing that these distances and angles could not be reliably determined by experiment. We conclude that DFT calculations are much better suited to determine the shape of the oxygen tetrahedra than X-ray or neutron diffraction on powders of moganite. Taking also the Si–O–Si angles into account, we found that among

the eight approaches considered in Table 1, the DMol³ code with the PBE functional gives the most convincing results for the bond angles in moganite.

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