BACTERIALLY INDUCED MINERALIZATION OF CALCIUM CARBONATE IN TERRESTRIAL ENVIRONMENTS: THE ROLE OF EXOPOLYSACCHARIDES AND AMINO ACIDS

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ABSTRACT: This study stresses the role of specific bacterial outer structures (such as glycocalix and parietal polymers) on calcium carbonate crystallization in terrestrial environments. The aim is to compare calcium carbonate crystals obtained in bacterial cultures with those obtained during abiotically mediated synthesis to show implications of exopolysaccharides and amino acids in the mineralogy and morphology of calcium carbonate crystals produced by living bacteria. This is done using various amounts of purified exopolysaccharide (xanthan EPS) and L-amino acids with a range of acidities. Amino acids and increasing xanthan content enhance sphere formation in calcite and vaterite. Regarding calcite, the morphology of crystals evolves from rhombohedral to needle shape. This evolution is characterized by stretching along the c axis as the amino acid changes from glutamine to aspartic acid and as the medium is progressively enriched in EPS. Regarding vaterite, the spherulitic habit is preserved throughout the morphological sequence and starts with spheres formed by the agglomeration of short needles, which are produced in a xanthan-free medium with glutamine. Monocrystals forming spheres increase in size as xanthan is added and the acidity of amino acids (glutamic and aspartic acids) is increased. At high xanthan concentrations, amino acids, and mainly aspartic and glutamic acids, induce vaterite precipitation. The role of the carboxyl group is also probably critical because bacterial outer structures associated with peptidoglycan commonly contain carboxyl groups. This role, combined with the results presented here, clearly demonstrate the influence of bacterial outer structure composition on the morphology and mineralogy of bacterially induced calcium carbonate. This point should not be neglected in the interpretation of calcite cements and carbonate accumulations in terrestrial environments.

INTRODUCTION

Microbiologically induced mineralization is defined as processes leading to inorganic mineral deposits by adventitious precipitation. This precipitation arises from secondary interactions between various metabolic processes producing carbonate species and the surrounding environment (Mann 2001). Mineral precipitation by microbes has been known about for a long time (see historical aspects in Ehrlich 1996, 1998). Among bacterially precipitated minerals, carbonates, and in particular calcium carbonate (CaCO₃) in numerous forms, are probably the most important (Ferris et al. 1989). Bacterially precipitated calcite (in terrestrial environments), and aragonite and high Mg calcite (in marine environments) have been described many times in the literature (e.g., Bouquet et al. 1973; Castanier et al. 2000; Chafetz 1986; Krumbein 1979). Although other species of calcium carbonate minerals are only rarely associated with bacterial activity, there are a couple of recently described examples: vaterite precipitation by the soil bacteria Xanthobacter autotrophicus (Braissant et al. 2002; Braissant and Verrecchia 2002) and monohydrocalcite (CaCO₃ · H₂O) by the halophilic bacteria Halomonas eurihalina (Rivadeneyra et al. 1998). The ecology of the microbes associated with vaterite and monohydrocalcite is consistent with the lacustrine environment in which the crystals formed (Giralt et al. 2001; Krumbein 1975) and not due to the activity of exotic strains under exceptional conditions. To our knowledge, microbially induced ikaite (CaCO₃ · 6H₂O) has not been reported, probably because of its instability.

FIG. 1.—X-ray diffractograms of CaCO₃ crystals associated with A) Xanthobacter autotrophicus and B) Ralstonia eutropha cultures. These diffractograms correspond perfectly to vaterite (A) and calcite (B), respectively.
Dolomite (CaMg(CO₃)₂) has been reported by Warthmann et al. (2000) to be precipitated by *Desulfovotovibrio* in anoxic marine environments.

Despite numerous reports of calcium carbonate precipitation by microbes and the important biomass they represent in soils, accumulation of terrestrial carbonates is still generally attributed to physicochemical processes (Lal et al. 2000). In this paper, it is demonstrated that relationships between morphologies and mineralogies of CaCO₃ encountered in soils and surficial sediments reveal a bacterial influence related to microbial biofilms. This study stresses the role of a specific exopolysaccharide (xanthan EPS) and amino acids on calcium carbonate crystallization in terrestrial environments.

In marine environments, outer structures such as the cyanobacterial S-layer have already been recognized as the main crystalline biorstructure able to act as a nucleus for calcium carbonate growth (see review in Schultze-Lam et al. 1992, 1996 and Smarda et al. 2002). Most microbial cells in natural environments form communities inside microbial biofilms (Decho 1990, 2000; Sutherland 2001 a, 2001 b). Therefore, exopolysaccharides (EPS) and amino acids most likely play an essential role in calcium carbonate morphology and mineralogy. This study investigates the potential relationships between EPS, amino acids, oxidative soil bacteria, and calcium carbonate crystals using conventional bacterial cultures and abiotically mediated calcium carbonate synthesis. The aim is to compare calcium carbonate crystals obtained in bacterial cultures with those obtained during abiotically mediated synthesis. This is done using various amounts of purified EPS and L-amino acids with a range of acidities and frequently associated with mucilages.

**MATERIALS AND METHODS**

Bacterially mediated calcium carbonate crystals were obtained after 20 days culturing of *Xanthobacter autotrophicus* (syn. *Corynebacterium autotrophicum*, DSM 432, ATCC 35674) and *Ralstonia eutropha* H16 (syn: *Alcaligenes eutrophi*, DSM: 428, ATCC 17699) on a B4 medium kept at 26°C (Merck yeast extract 4.0 g/L; Merck calcium acetate 2.5 g/L; Merck agar-agar 15 g/L; Boquet et al. 1973). These two strains are oxidative terrestrial bacteria commonly found in soils and sediments. In the cultures, the main difference between the two strains is the amount of polysaccharide produced: *X. autotrophicus* is surrounded by a large amount of slime, whereas *R. eutropha* produces only an extremely low amount of glycoalix (Holt et al. 1993). The pH was not monitored during culturing, but final pH was > 9 as revealed by a droplet of pH indicator. Unincoculated Petri dishes were kept as a sterility control.

Abiotically mediated crystals have been obtained by precipitation from a CaCl₂ / (NH₄)₂CO₃ reaction inside a polysaccharidic medium (EPS) with various types of L-amino acids. The aim of this experiment is to simulate precipitation of calcium carbonate in an EPS-rich environment without involving living bacteria. Ten runs of this experiment have been conducted. The synthesis was performed using a CaCl₂ 0.05 M solution to which was added: (1) 1.0% (w/v) amino acids (L-glutamine, L-glutamic acid and L-aspartic acid). These amino acids were chosen because they are commonly found in the parietal structures of many bacteria, e.g. polyglutamate in *Bacillus anthracis*, *Mycobacterium tuberculosis*, *Sporosarcina halophila*, polyglutamine in the genus *Xanthobacter* and *Flexithrix*, and finally polylaspartate in *Synechococcus* sp.); and (2) 0%, 0.1%, 0.5% and 1.0% (w/v) concentrations of xanthan from Sigma™ (a bacterial EPS mainly composed of mannose and glucose). The pH of this solution was adjusted to 8.4 (the pH of calcite stability at 1 atm and 25°C) before adding xanthan. The solution was sterilized by autoclaving at 121°C for one hour. After sterilization, the solutions were placed in Petri dishes that remained for 20 days in ethanol-washed desiccators filled with (NH₄)₂CO₃. Petri dishes were examined with a binocular microscope to control possible presence of contaminants.

Crystals obtained in bacterial cultures, as well as in the abiotic experiment, were both isolated by either sedimentation or directly from the culture using precision tweezers. The crystals obtained were gold coated and observed with a Philips XL20 SEM. Crystals sampled from bacterial colonies were previously washed in saturated calcium hypochlorite solution in order to remove organic matter from the surfaces of the crystals. Samples of bacterial colonies bearing crystals were fixed with glutaraldehyde (5%), immediately dehydrated in ethanol, and air-dried after ethanol replacement by tetramethylsilane (TMS; Dey et al. 1989). These samples were gold sputter coated for 60 seconds to give a 23 nm gold coating, and observed with a Philips XL20 SEM. Crystals were analyzed by X-ray diffraction (XRD) using a Scintag diffractometer and by an energy-dispersive spectrometry (EDS) microprobe coupled with a Philips XL30 environmental scanning electron microscope (ESEM). Calcite and vaterite mineralogies were determined by cross checking XRD analyses, crystal habits, and relative proportions of each phase plus comparison of observed crystal shapes with shapes documented in the literature.

**RESULTS**

*Calcium Carbonate Crystals Obtained in Bacterial Cultures*

Bacterially mediated crystals were found mostly on the surfaces of colonies. They have various mineralogies and morphologies. In terms of mineralogy, *X. autotrophicus* produced mainly brown vaterite spherulites ranging from 50 μm to 200 μm in diameter, with a small fraction of calcite crystals (Fig. 1A), whereas *R. eutropha* produced colorless calcite crystals ranging from 100 μm to 600 μm in size, with traces of vaterite (Fig. 1B).

Spheres produced by *X. autotrophicus* show three different monocrysyl (small individual crystal) arrangements (Fig. 2A–D). Although the overall shape of a crystal remains, the morphology of individual crystals forming the sphere varies from a fan shape to a needle (Fig. 3). The crystals are organized around a central point and grow radially to form the sphere. Accessory shapes are also associated with the spheres. They are made of either a polyhedral assemblage of calcite crystals or a mixture of calcite and vaterite. Two main morphologies have been observed. “Fried egg” shapes are probably formed at the surface of the colony (Fig. 2E–G). The sub-spherical part is constituted by vaterite and grows on a lens- to plate-shape calcite surface. The flat part is in contact with the atmosphere, the sphere being inside the colony. Therefore, the general shape is a flipped “fried egg.” Calcite also forms imperfect spherical clusters made of imbricated rhombohedra (Fig. 2H).

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**Fig. 2.—Morphology of crystals obtained during bacterial growth. A–H. *Xanthobacter autotrophicus* ; I–P. *Ralstonia eutropha*. A) General view of vaterite spherulites in cross-polarized light (XPL). B) Close-up of part A showing two different crystal arrangements. Spheres showing a black cross are constituted by needles, whereas other spheres are formed by clusters of small monocrystals. C, D) Scanning electron microscope (SEM) view of individual (C) and coalescent (D) vaterite spherulites. The surface of the sphere is not smooth and shows a mammillated rough topography. E, F) “Fried egg” morphology of accessory crystals. These clusters are constituted by two parts, a calcitic bottom part having a lens- to plate-shape and a central hemispherical upper part constituted by vaterite crystals of various sizes. G) Detail of the hemispherical part showing small hexagons of vaterite. H) Imbricated rhombohedra forming imperfect spherical calcite clusters. I) Subhedral to euhedral rhombohedra. XPL view. J) ESEM view of a calcite flower formed by flat sparitic crystals. K) SEM view of the heart of a calcite flower constituted by crystals forming a spiral-like structure. L) SEM view of a calcite flower with petals formed by thin palisadic crystals. M) XPL view of a calcite cluster constituted by four triangular-shaped crystals joined at their tip to form a thickened Maltese cross. N) SEM view of a similar crystal shown in part M. O, P) XPL and SEM view of funnel-like crystals.
Crystals produced by *R. eutropha* have three main shapes. The two most common morphologies encountered are euhedral to subhedral rhombohedra (Fig. 2I) and calcite flowers, composed of crystals with euhedral terminations (Fig. 2J–L). The center of the flower is formed by a circular spiral-like crystal (Fig. 2K). Two other shapes have been observed: thickened “Maltese crosses” (Fig. 2M, N) and, rarely funnels (Fig. 2O, P). In blank uninoculated Petri dishes, crystals were not observed, emphasizing the crucial role of bacteria.

**Calcium Carbonate Crystals Obtained in Abiotically Mediated Experiments**

The reaction between CaCl₂ and (NH₄)₂CO₃, with neither xanthan nor amino acids present, produces calcite rhombohedra (Fig. 4A), which is not surprising. However, synthesis of calcium carbonate crystals in mucilaginous solutions results in numerous crystal morphologies, emphasizing the potential role of EPS in their shape diversity. Synthesis using increasing concentrations of xanthan without the addition of amino acids leads mainly to individual and branching crystals. At a low concentration of xanthan (0.1%), dendritic crystals form. Dendrites can be disordered, and similar to structures obtained by diffusion-limited aggregation (Fig. 4B), or ordered along the main three directions of calcite crystal growth (Fig. 4B). Nevertheless, subhedral rhombohedra remain the most abundant shape (Fig. 4B). At a higher concentration of xanthan (0.5%), the most common shape (constituting about 90%) is a dendrite-like crystal, which is in fact lacunar and made of small stacked rhombohedra (Fig. 4C). Around 5% of the crystals are euhedral rhombohedra (Fig. 4C). The remaining 5% are imbricated rhombohedron twins that form irregular spheroidal clusters (Fig. 4C). When the xanthan concentration reaches 1%, more and more imbricated rhombohedron twins agglomerate until they form subspherical clusters. Whatever the concentration in xanthan, all the crystals obtained in the amino acid-free media are calcitic.

When amino acids are added to the medium, various crystal morphologies are obtained (Fig. 4E–P). At first glance, it is difficult to clearly see a trend in morphologies, although the habit becomes more spherical as xanthan increases. In addition, the two calcium carbonate morphologies—vaterite and calcite—are always precipitated during abiotic synthesis in the presence of amino acids. Nevertheless, the ratio between the two mineralogical species seems to be essentially related to the abundance of EPS and the nature of amino acids. Aspartic acid, glutamic acid, and glutamine always induced vaterite formation at a high concentration of xanthan (1.0%), although calcite is still present. Consequently, it is critical to discriminate between the two mineralogies in order to understand the difference in morphologies in Figure 4. Two morphological sequences (Fig. 5) can be described in relation to increasing EPS content (which acts as a weak acid) and the acidity of the added amino acid. Glutamine is considered as a basic amino acid, whereas glutamic and aspartic acids are acidic, aspartic acid being more acidic than glutamic acid.

Regarding calcite (Fig. 5), the morphological sequence starts with rhombohedra produced in a xanthan-free medium with glutamine (Fig. 4E). The conditions in this case are the most basic (highest pH) of the total experiment involving amino acids. By adding xanthan and by increasing the acidity of amino acids (glutamic and aspartic acids), rhombohedron edges become smoother (Fig. 4G) and crystals form numerous twins, agglomerated around a center (Fig. 4G, H), and finally evolve to fibro-radial spheres (Fig. 4K–P). This sequence can be summarized by the decrease of the monocrystal size forming the shape, from calcite rhombohedra to calcite styloids followed by needles (Fig. 5).

Regarding vaterite (Fig. 5), the spherulitic habit is preserved throughout the morphological sequence. It starts with spheres formed by the agglomeration of short needles, which are produced in a xanthan-free glutamine medium (Fig. 4E). Nevertheless, some spheres are composed of larger (but still small) monocrystals (Fig. 4E). These monocrystals increase in size as xanthan is added and the acidity of amino acids (glutamic and aspartic acids) is increased (Fig. 4I, J, L–P). Contrary to the sequence described for calcite, the vaterite sequence can be illustrated by the increase in the size of the monocrystals forming the spheres. Nevertheless, in the presence of glutamine and glutamic acid and a high amount of EPS, vaterite can form fibro-radial spherulites, which show a black cross in polarized light, as shown by Dedek (1966). The only way to differentiate calcite from this vaterite is by using XRD.

Figure 4—Scanning electron microscope photographs of crystal morphologies obtained during abiotically mediated synthesis of calcium carbonate in the presence of exopolysaccharides and amino acids. Abscissa: Blank (no amino acids present), Gln, L-glutamine; Glu, L-glutamic acid, Asp, L-aspartic acid. Coordinates: xanthan content (a glucose and mannose polymer), from 0.0% (absence of xanthan) to 0.1, 0.5 and 1.0% w/v. A) Euhedral calcite crystals. B) Subhedral and dendritic calcite crystals. Dendritic crystals show two different morphologies: diffusion-limited aggregation clusters and ordered dendrites. C) Crystal morphologies become chunkier, leading to imbricated twins. Some euhedral rhombohedra are preserved. D) Agglomerated twins tending to form spheres. E) Euhedral calcite rhombohedra associated with vaterite spherulites composed of either short needle monocrystals or an agglomerate of small euhedral crystals. F) Imperfect calcite rhombohedra are present with vaterite spheres. General structure of vaterite spherulites is similar to part E. G) Rare spherulitic vaterite associated with calcite rhombohedra. Edges of calcite crystals are smoother than in part A, and they tend to form imbricated twin clusters. H) Vaterite and calcite spherulites. Calcite appears as imbricated twin clusters. In addition, both vaterite and calcite can be characterized by fibro-radial spheres. I) Upper part: vaterite spiky agglomerates and cauliflower-shaped calcite. Bottom part: left, calcite rhombohedra; right, epitactic growth of a hexagonal vaterite crystal on a calcite substrate forming cauliflower clusters. J) Spiky agglomerate of vaterite starting to form a sphere, and a rough calcite spherulite. K) Calcite spheres. The arrow shows remains of a rhombohedron emphasizing the transition between imbricated twin clusters shown in parts G and H and the structure of spheres formed by styloidal crystals. Double sphere of calcite showing a structure close to fibro-radial spheres and composed by styloidal monocrystals. L) Fibro-radial calcite and vaterite spheres with a smooth surface associated with spiky vaterite spherulites (arrow). M) Fibro-radial calcite spherulites (right). Spiky vaterite agglomerate (left). N) General view of spiky vaterite spherulites associated with calcite spheres. Flat shapes are due to contact of the surface of the medium with the atmosphere. O) Calcite spheres. Some spheres are constituted by stacked flat monocrystals, which can be compared with spheres formed by styloidal crystals shown in part K. P) General view of vaterite and calcite spheres. Vaterite spherulites are characterized by the agglomeration of large monocrystals, whereas calcite spheres are smooth and fibro-radial.
Fig. 5.—Sequences of calcite and vaterite morphologies obtained during the abiotic experiment. This sketch shows the morphologies related to increasing amino acid acidity and xanthan content. Scales are relative. The various domains of influence of amino acids are shown. Six main steps can be described for both vaterite and calcite. The calcite sequence starts with rhombohedra. Size of monocrystals forming the clusters decreases from Step 1 to Step 6 until they form needles. The general shape of calcite evolves from rhombohedra to fibro-radial spherulites. Vaterite sequence: vaterite is always precipitated as spheres. The monocrystals constituting the spheres increase in size from Step 1 (short needles) to Step 6 (large hexagons). Between Step 2 and Step 3, vaterite can occur as fibro-radial spherulites. Sketches were traced from photographs.

In summary, amino acids and increasing xanthan content enhance sphere formation in both calcite and vaterite. In addition, regarding calcite, the morphology of rhombohedra evolves from a euhedral shape to a needle. This evolution is characterized by stretching along the $c$ axis as the amino acid changes from glutamine to aspartic acid and as the medium is progressively enriched in EPS.

DISCUSSION

Many spherulitic features are observed in carbonate soils and surficial sediments. The explanation of their origin is often unclear and usually attributed to possible organic influence without further elucidation. Spherulitic or oolitic habit have been observed in many other cases involving different microorganisms and different mineral species (e.g., Folk 1993). Calcium carbonate spherulites remain the most commonly described features. It has been shown that formation of aragonitic spherulites by Deleya halophila may be the result of cell calcification and aggregation (Rivadeneyra et al. 1996). In addition, formation of magnesian calcite spherulites and dumbbells has been related to the slime-producing bacteria, Myxococcus xanthus (González-Muñoz et al. 2000; Holt et al. 1993). Various types of apatite $\left[\text{Ca}_{10}(\text{PO}_4)_{6-x}(\text{CO}_3)_x(\text{F},\text{OH})_{2+x}\right]$ have been shown to precipitate with a spherulitic habit when under the influence of organic matter or during (nano)bacterial activity. Kajander and Çifşanglu (1998) have shown that nanobacteria can be responsible for the formation of small spherulites (about 2 $\mu$m in diameter) alone or inside human cells (the spherulitic habit is not surprising considering the high density of macromolecules). Apatite spherulites including proteins (referred to as nanoforms) have also been shown to precipitate in a medium containing sterile fetal bovine serum (Vali et al. 2001). Salt ooids have been attributed to halophilic bacterial activity, which may guide the growth of biominerals (Castanier et al. 1999).

In conclusion, recent research illustrates the numerous possible interactions between crystals and organic matter. Our results emphasize the importance of bacterial biofilms containing exopolysaccharides and amino acids in the precipitation of $\text{CaCO}_3$, and resemble features encountered in bacterial cultures, soils, and paleosols.

The morphologies of crystals obtained in $X$. autotrophicus and $R$. eutrophus cultures (Fig. 2) can now be understood given the results depicted in Figure 4. Xanthobacter autotrophicus is known to produce substantial amounts of exopolysaccharides (Wiegel 1991). This property leads to the precipitation of vaterite spherulites (Fig. 2A–D), with some spheres showing a black cross of extinction in cross-polarized light. These morphologies are equivalent to those observed in Step 3 of Figure 5. Vaterite spherulites showing a black cross have also been observed, although rarely in the presence of glutamine with 0.5% xanthan. Regarding calcite, the accessory shape constituted by imbricated rhombohedra (Fig. 2H) is also equivalent to Step 3 of Figure 5. These observations indicate that crystals observed in the $X$. autotrophicus culture should be produced in a medium enriched
in EPS as well as glutamine. Indeed, *X. autotrophicus* makes not only large amounts of exopolysaccharides but also a polyglutamine parietal polymer (Kandler et al. 1983; Wiegel 1991).

In contrast, the fact that *R. eutropha* only produces calcite rhombohedra and flower-shaped agglomerates is linked to an absence or only 0.1% EPS in the medium. The calcite crystals can be compared with Step 1 of calcite in Figure 5 or with crystals shown in Figures 4E and F. This is consistent with the fact that *R. eutropha* produces only low amounts of EPS.

The results obtained in Figure 4 can also be compared with features observed in natural environments. Spherulites showing an extinction cross in cross-polarized light are commonly found in soils and paleosols (Fig. 6). These spherulites infill pores and are associated with micritic and/or organo-micritic layers (Fig. 6A, B). Their origin has never been clearly identified. Nevertheless, their morphology (Fig. 6C, D) is identical to crystals produced in the EPS-rich medium (Fig. 4H, K, L, P). This point indicates the influence of biofilms in the precipitation of such spherulites in soils. A similar calcium carbonate spherulite has also been observed in a polysaccharide-rich tropical soil (Fig. 7A). Calcite and vaterite spherulites have also been reported in the mucilagenous sheath of cyanobacteria, indicating the fundamental role exopolysaccharides can play in their precipitation (Verrecchia et al. 1995; Giralt et al. 2001). In a different tropical soil, a feature identical to those shown in Figures 4J, M, and N has also been observed: it is a vaterite cluster growing on a quartz grain in a biofilm rich in EPS and amino acids (Fig. 7B). In conclusion, the abiotic experiments conducted in this study shed new light on the potential contribution of bacterial EPS and amino acids on calcium carbonate precipitation in terrestrial environments.

In natural environments, the variation of CaCO$_3$ crystal morphologies and mineralogies associated with EPS may result from its heterogeneity due to the diversity of microorganisms living within a single biofilm. Heterogeneity in slimes leads to differences in diffusion coefficients (local supersaturation), surface tensions, and viscosity. The viscosity of the 1.0%
CONCLUSIONS

From the results presented in this paper, it seems likely that spheralitic calcite and/or vaterite crystals may indicate the presence of mucilaginous bacteria or biofilms at the time of formation (and probably other mucilage-producing organisms, e.g., cyanobacteria or fungi). The relative abundance of mineral species (calcite vs. vaterite) can be informative in determining the nature of the biofilm, i.e., if they contain polyspartic acid or other polyamino acids. In terms of sedimentary petrology, these results question the origin of some cements found in terrestrial sediments that were previously attributed to purely physicochemical processes.

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