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Eocene-Pliocene time scale and stratigraphy of the Upper Rhine Graben (URG) and the Swiss Molasse Basin (SMB)

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Abstract We present a general stratigraphic synthesis for the Upper Rhine Graben (URG) and the Swiss Molasse Basin (SMB) from Eocene to Pliocene times. The stratigraphic data were compiled both from literature and from research carried out by the authors during the past 6 years; an index of the stratigraphically most important localities is provided. We distinguish 14 geographical areas from the Helvetic domain in the South to the Hanau Basin in the North. For each geographical area, we give a synthesis of the biostratigraphy, lithofacies, and chronostratigraphic ranges. The relationships between this stratigraphic record and the global sea-level changes are generally disturbed by the geodynamic (e.g., subsidence) evolution of the basins. However, global sea-level changes probably affected the dynamic of transgression–regression in the URG (e.g., Middle Pechelbronn Beds and Serie Grise corresponding with sea-level rise between Ru1/Ru2 and Ru2/Ru3 sequences, respectively) as well as in the Molasse basin (regression of the UMM corresponding with the sea-level drop at the Ch1 sequence). The URGENT-project (Upper Rhine Graben evolution and

neotectonics) provided an unique opportunity to carry out and present this synthesis. Discussions with scientists addressing sedimentology, tectonics, geophysics and geochemistry permitted the comparison of the sedimentary history and stratigraphy of the basin with processes controlling its geodynamic evolution. Data presented here back up the palaeogeographic reconstructions presented in a companion paper by the same authors (see Berger et al. in Int J Earth Sci 2005).

Keywords Rhine Graben · Molasse · Paleogene · Neogene · Stratigraphy

Introduction

During the last decade, important progresses have been made in Tertiary stratigraphy and correlation. Different synthetic stratigraphic charts have been published (e.g., Berggren et al. 1995; Hardenbol et al. 1998; Steininger 1999). Some recent data slightly modify several boundaries (particularly owing to the extension of astronomically calibrated geological time scales into the Neogene and Paleogene; see Pälike and Shackleton 2003) that will be synthesized probably in 2005 (see Berggren et al. 2003).

A detailed stratigraphic frame is absolutely necessary to constrain geodynamic models. This stratigraphic study could thus be used to evaluate different geodynamic questions such as subsidence rates, uplift of Vogesen–Schwarzwald massifs or correlations between Upper Rhine Graben (URG)-rifting and alpine orogeny.

The charts presented here, for the URG and Western Switzerland, address two basic aspects:

A. The first aspect concerns the time scale, correlating biozones, magnetostratigraphy, sequence stratigraphy, chronostratigraphic units and the absolute time scale (Fig. 1). We present in this chart different correlations proposed by different authors, following the concept

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published by Berger (1992a). Eight columns have been compiled by the following authors:

The correlation presented here has to be used care-

3. Plateau Molasse distal

4. Jura Molasse SW (Valserine, Joux, Auberson, Travers, Val de Ruz)

- Magnetostratigraphy:	Cande and Kent 1992, 1995
- Planktonic foram. zones:	1. Berggren et al. 1995 2. Hardenbol et al. 1998 3. Steininger 1999
- Calc. Nannoplankton zones:	1. Berggren et al. 1995 2. Hardenbol et al. 1998 3. Steininger 1999 4. Odin et al. 1997
- Mammal zones and Swiss mammal levels:	2. Hardenbol et al. 1998 3. Steininger 1999 5. Schmidt-kittler et al. 1997 6. Bolliger 1997, Engesser and Mödden 1997, Kälín 1997a, Kälín and Kempf 2002, Kempf et al. 1997, 1999, Kempf and Matter 1999, Schlunegger et al. 1993, 1996, 1997a,b, c 7. Steininger et al. 1996 8. Heissig 1997 9. Fejfar et al. 1997
- Charophyte zones:	10. Escarguel et al. 1997, Legendre and Leveque 1997 2. Hardenbol et al. 1998 11. Riveline et al. 1996 12. Berger 1999
- Otolith zones:	13. Reichenbacher 1999, 2000
- Sequence stratigraphy:	2. Hardenbol et al. 1998 3. Steininger 1999
- Paratethys stages:	14. Vakarcs et al. 1998
- Mediterranean stages:	1. Berggren et al. 1995 2. Hardenbol et al. 1998 7. Steininger et al. 1996 15. Odin and Luterbacher 1992

Accepted GSSP according to Aguirre and Pasini (1985), Castradori et al. (1998), Hilgen et al. (2000), Premoli Silva and Jenkins (1993), Remane et al. (1999), Rio et al. (1998), Steininger et al. (1997), and van Couvering et al. (1999)
Possible future GSSP according to Odin et al. (1997)

fully as demonstrated in Fig. 2: for instance, if a locality A (continental) is dated by mammals as MP21, and a locality B (marine) is dated by nannoplankton as NP22, the locality A could be older, synchronous or younger than the locality B due to the following factors:

- The boundaries between NP21, NP22 and NP23 are generally based on extinction events. Thus, serious problems may appear due to reworked nannofossils. Therefore, correlations between these boundaries and the absolute time scale are still uncertain.
- The precise age of the boundary between MP21 and MP22 varies from about 33.0 Ma to 32.0 Ma., according to different authors.

These examples explain the difficulties that often appear in correlating different environments (marine, continental). They underline that the quality of the different data has to be checked carefully, before using for other purposes such as paleogeography or paleoclimate changes.

B. The second aspect of our chart presents the lithostratigraphic units of the Swiss Molasse Basin (SMB) and the URG (Figs. 7, 8, 9, 10). The 14 columns for the specific geographical areas (Fig. 3) are listed below :

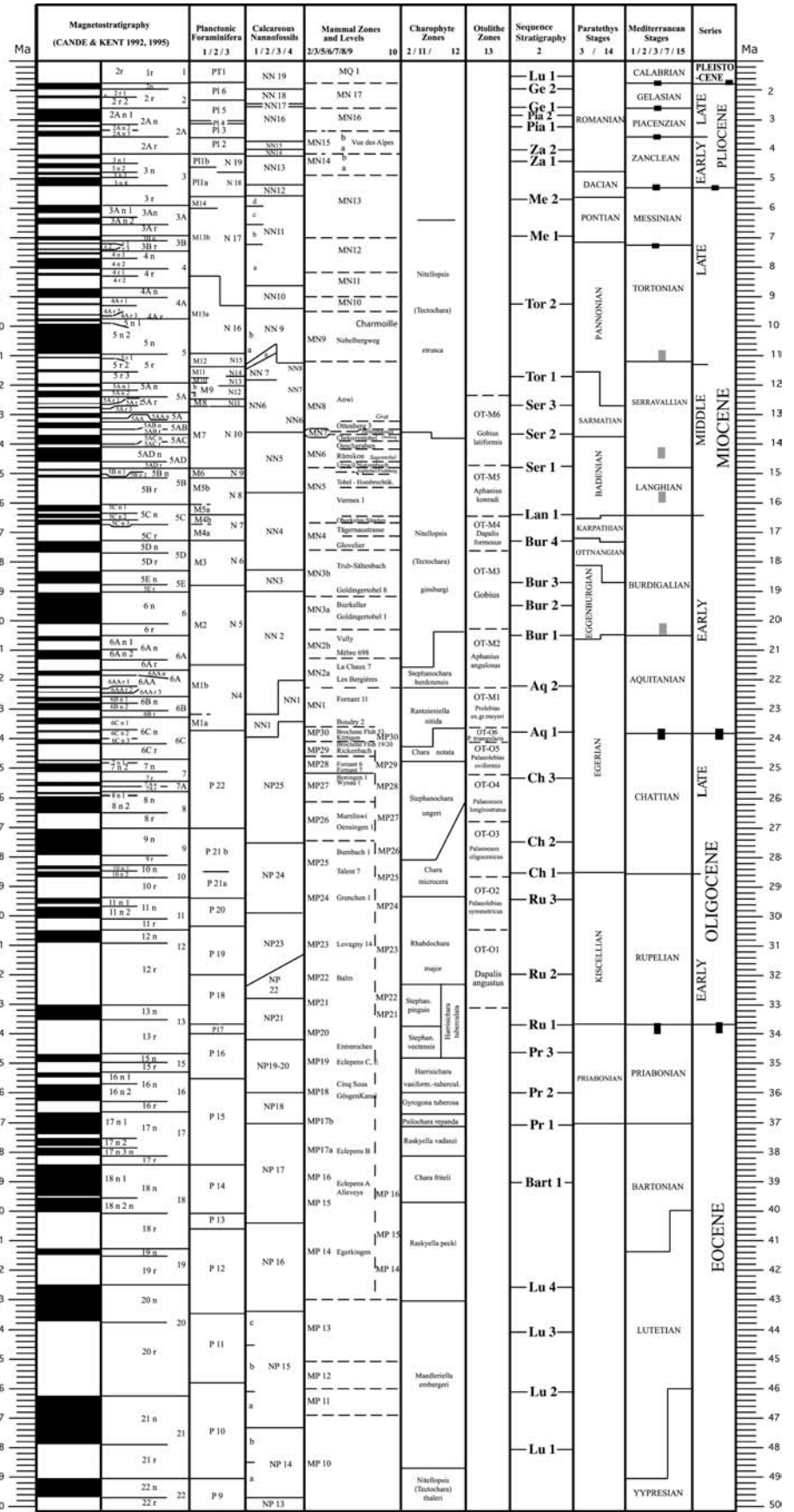
1. Helvetic domain and Subalpine Molasse
2. Plateau Molasse proximal

5. Jura Molasse NW (Verrières, Pont de Martel, Locle-Chaux de Fonds)
6. Jura Molasse Centre-S (St. Imier, Pery-Reuchenette, Tramelan-Tavannes, Welschenrohr-Balstahl, Sornetan-Bellelay, Moutier)
7. Jura Molasse Centre-N (Soulce, Delémont, Laufen, Courgenay-Charmoilles, Bressaucourt-Porrentruy, Liesberg, Löwenburg, Movelier)
8. Jura Molasse E (Mummliwil, Waldenburg)
9. Southern URG (Basel Horst, Dannemarie Basin, Mulhouse Horst, Mulhouse Basin potassique, Sierentz-Wollschwiler Basin, Rauracian Depression, Lörrach to Freiburg im Br.)
10. S-Middle URG (Colmar, Sélestat, Erstein, Zorn or Strasbourg Basin), Fault area of Saverne, Ribeauvillé & Guebwiller
11. N-Middle URG (Hagenau, Pechelbronn, Rastatt, Karlsruhe, Landau), Fault area of Bruchsal-Wiesloch
12. Northern URG (Heidelberg Basin, Darmstadt, Wetterau, Horloff Graben)

S-boundary: Neustadt/Weinstraße-Heidelberg / W-boundary: Mainz Basin / E-boundary: Hanau Basin

13. Mainz Basin
14. Hanau Basin

Fig. 1 Eocene-Pliocene stratigraphic chart



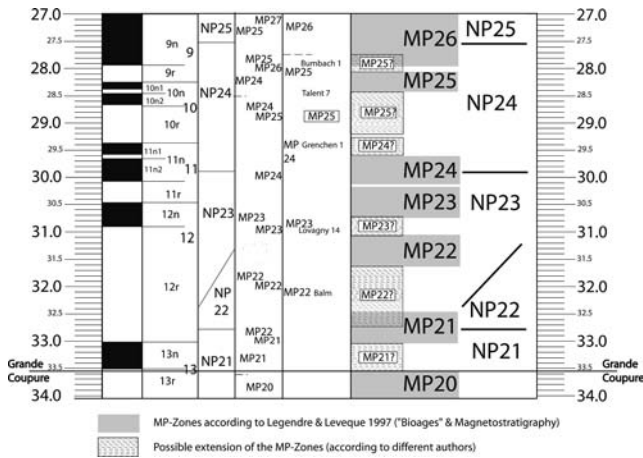


Fig. 2 Early Oligocene correlation problems between mammals (MP21–25) and calcareous nannoplankton (NP21–25)

The lithostratigraphic units are correlated with the time scale according to stratigraphical data found in outcrops or boreholes. Figures 4 and 5 list the most important localities, which we used for stratigraphic correlations. Figure 6 summarizes the lithostratigraphical units (and their key numbers), which we applied in

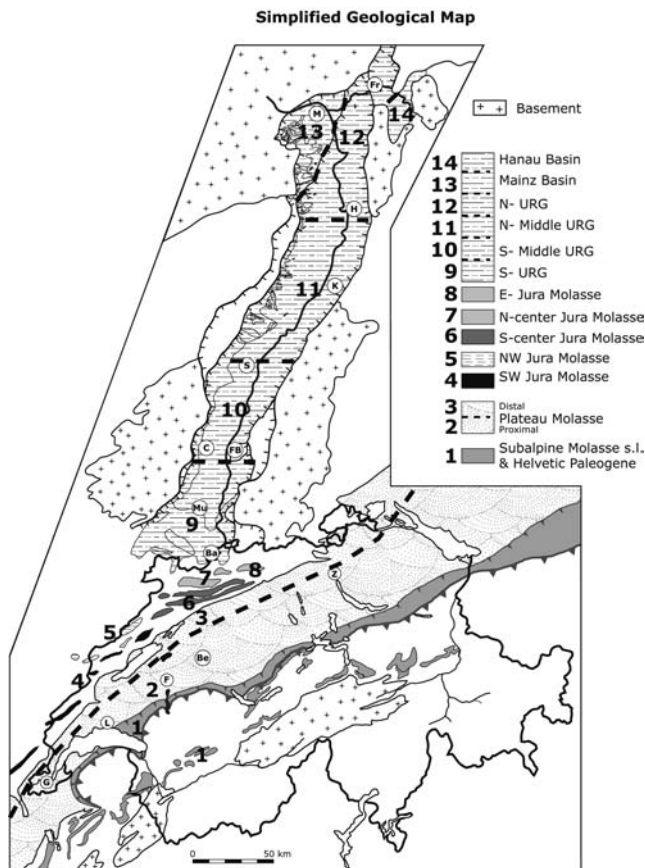


Fig. 3 Simplified geological map

Fig. 4 Map of localities (see listing on Fig. 5)

Figs. 7, 8, 9 and 10. To avoid the abundant use of quotation marks in the following discussion, informal lithostratigraphic units, like *Green marls* or *Calcaires et dolomies*, are written in italics.

The Deutsche Stratigraphische Kommission (2002) recently published a stratigraphic synthesis for the German part of the URG. The study of Sissingh (2003) deals with the Rhine Graben system in general, and stresses the correlation of eustatic sea-level changes, rifting phases, palaeogeography, and sedimentary evolution. However, the topic of our study is to present both the stratigraphical frame and the sedimentary evolution for the URG–Western Switzerland system. For this purpose, we took into account a lot of new stratigraphic data.

The relationships between the stratigraphic record presented here and the global sea-level changes are generally disturbed by the geodynamic (subsidence) evolution of the basins. However, global sea-level changes probably affected the dynamic of transgression–regression in the URG as well as in the Molasse basin (see discussion below).

Helvetic domain subalpine Molasse and Foreland Molasse (Fig. 7)

Eocene

Deep marine sediments of Lutetian to Priabonian age are present in the Helvetic nappes (Einsiedeln Formation, Steinbach-Gallensis Fm., Bürgen Fm., *Globigerinenmergel*, *Flysch sudhelvétique*, *Marnes à Foraminifères*, Hohgant Fm.). In addition, coastal facies (Klimsenhorn Fm., Wildstrubel Fm.) and brackish environments (*Couches à Cerithes*, *Couches des Diablerets*, Sanetsch Fm.) of the same age are known (Herb 1988; Herb et al. 1984; Menkveld–Gfeller 1994, 1995, 1997; Weidmann et al. 1991).

In the Plateau Molasse, continental deposits, rich in lateritic remains (the so-called *Siderolithic*, with *Bohnerz* and/or *Hupper* = quartz sands), were deposited, some of them dated by mammals (Egerkingen, Eclepens, Gösgen Kanal, see Engesser and Mayo 1987; Engesser and Mödden 1997; Hooker and Weidmann 2000).

Oligocene

In the Helvetic domain and the subalpine Molasse, marine sedimentation continued with the Lower Marine Molasse (UMM) during the Rupelian. The UMM comprises the *Meletta-shales*, the *Taveyannaz* and *Aldorf sandstones*, the Val d'Illeuz Fm. and the Vaulruz Fm. (see

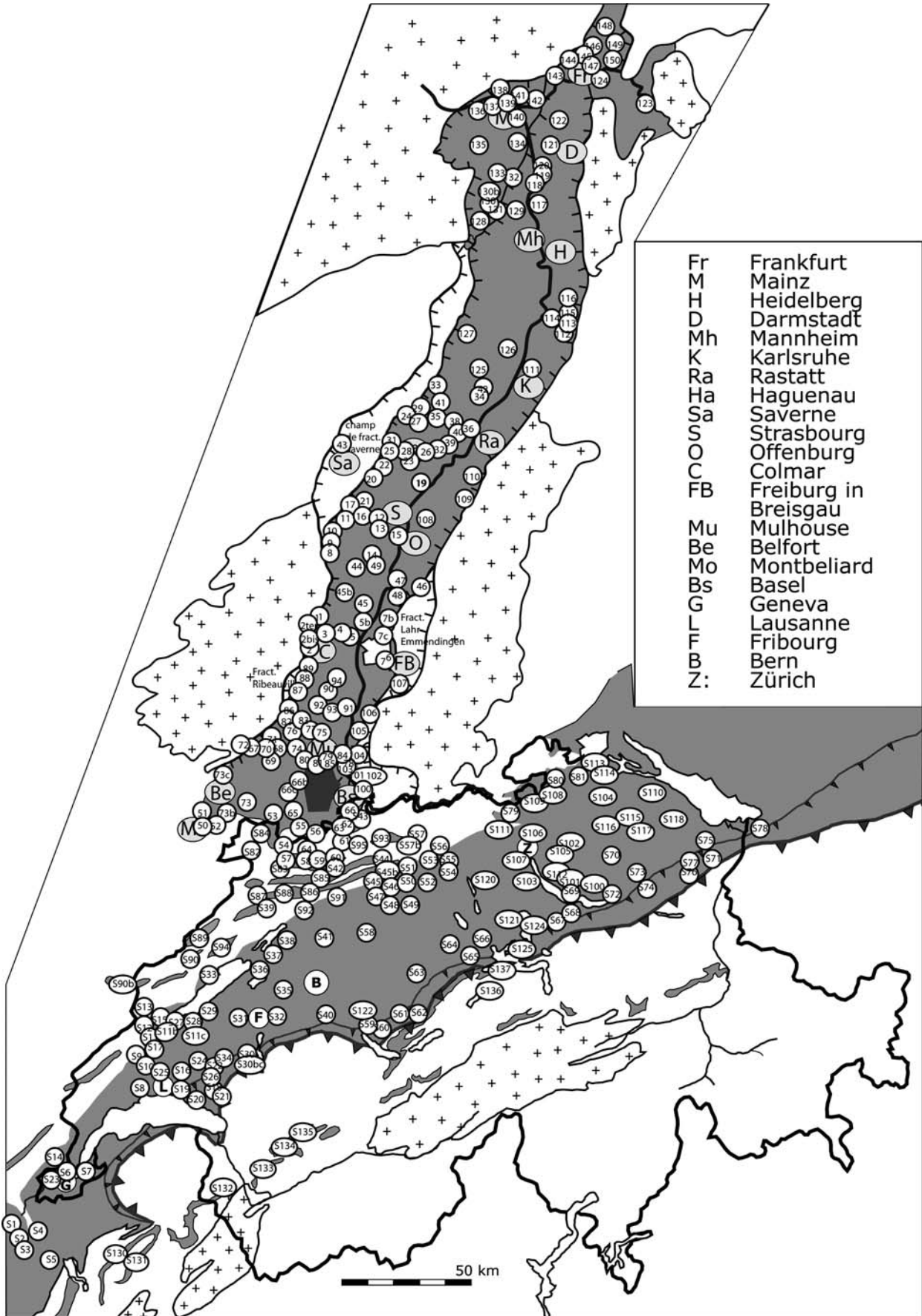


Fig. 5 List of the localities (see map on Fig. 4)

1. Rohrschwir	100. Tüllingen	S47. Bannwil
2. Sigolsheim	101. Istein, Kelinkems	S48. Wischberg
2bis Turkheim, Wintzenheim,	102. Hammerstein, Wolzen,	S49. Roggliswil
Wettolsheim	Kandern, Röttein	S50. Boningen
2ter Bergheim-Mittelwahr	103. Sierentz	S51. Rickenbach
3. Ostheim	104. Mülheim-Schiengen	S52. Bierkeller, Brunngraben, hinterer Bühnenberg, Safenwil
4. Illhausen	105. Oberweiler, Zienken,	S53. Gösgen-Kanal
5. Eisenheim	Bitzingen, Stauffen	S54. Hirschtal
5b. Oehnenheim	106. Grissheim	S55. Gränichen-Moorberg
6. Eichstätt-Pfaffenthal-	107. Pfaffenweiler-Merzhausen	S56. Küttigen
Bölzingen	108. Kork	S57. Anwil
7. Wasenweiler	109. Bühl 1	S57b. Zeglingen
7b. Herbolzheim, Emmendingen	110. Baden-Baden 1 & 2	S58. Wyringen-Brittenberg
7c. Limberg	111. Linkenheim 1	S59. Schwendibach, Frohberg, Tobel-Hombrechtikon,
8. Heiligenstein	112. Ubadst	Hottwil, Chlaustobel, Matt Tébeli, Martinsbrunneli,
9. Obernai, Bischofsheim, Barr	113. Malsch sondages	Tägernastrasse, Hummelberg
10. Bischofsheim-Rosheim	114. Rott-Malsch	S60. Loseneegg 1+2, Marbachsection, Cheilstisteig,
11. Ergersheim-Scharachberg-Wolkheim	115. Oberhausen 1	Dürenbach, Pressenbach, Thun 1, Eriz
12. Entzheim-Hölzrheim	116. Wiesloch, Rauenberg	S61. Hornbach 3
13. Lipsheim	117. Hofheim 3	S62. Bumbach 1, Emme Section, Trub-Sältenbach
14. Schaeffersheim	118. Eich 17	S63. Hasenbach 1, Eimattli, Oeschgraben,
15. Eschau	119. Biebesheim 2	Schwändigraben, Fontannen Section
16. Kolbsheim	120. Stockstadt	S64. Werthenstein-Grabenhüsil
17. Ittlenheim	121. Dornheim 9	S65. Rümliq 2, Ränggloch 3/8
18. Truchtersheim-Kelinfrankenheim	122. Königsstädten	S66. Löwencenter, Hintersteinbruch, Luzern 10/76
-Reitwiler	123. Gross Anhaime-	S67. Sparenweid
19. Gamsheim-Killstett	Gross Katzenburg	S68. Feusisberg
20. Kienheim	124. Obertschhausen	S69. Schillfitobel, Bürgistobel B, Tobel
21. Dossenheim-Wiwenheim-	125. Wnden 4	+Speerstrasse Hombrechtikon, Matt
Quatzenheim	126. Rülzheim 1 & 2	S70. Ergeten, Steg, Chlihörnl, Grat,
22. Schaffhouse-Hochfelden	127. Landau west 1	Bärtobel-Hörnli, Goggelswald, Gfell
23. Bischwiller	128. Gölheim 1, Kerzenheim,	S71. Sommersberg
24. Woerth (+ Diffenbach)	Lautersheim 1	S72. Goldingertobel, Lattenbach, Fätzikon
25. Wittersheim	129. Worms 3	S73. Dorfbachtobel, Wattwil
26. Schirrheim	130. Alzey 1, 2, Heimersheim, Weinheim	S74. Ebnat-Kappel, Wintersberg-Trempe,
27. Schwabwiler	130b. Uffhoben, Flonheim	Umfahrungsstrasse+Thur-Krummenau
28. Schweighouse (2) & Ohlungen	131. Galgenberg, Eppelsheim,	Steintal section
29. Merkwiler - Pechelbronn, Lobsann	Westhofen, Gundersheim,	S75. Martinsbrücke
+ Soultz s- Forets & Retschwiler	Bermersheim, Esselborn	S76. Appenzel-Kaubach
30. Schwabwiler	132. Alsheim, Dorn-Dürkheim	S77. Säglbach
31. Dauendorf	133. Gabsheim	S78. Unter Staudach, Schwarzachtobel
32. Soufflenheim-Sessenheim	134. Oppenheim 1, Nierstein, Bodenheim	S79. Kaiserstuhl
33. Altenstadt-Oberhof, Riedselz,	135. Sprendlingen 3, Wissberg 2,	S80. Benken am Kohlfirst
Wissenbourg	Wolfsheim	S81. Schlätlingen
34. Scheibenhard	136. Algesheim, Ingelheim	S82. Bressaucourt
35. Rittershofen	137. Heidesheim	S83. Glovelier, Neuf Camps
36. Beinheim	138. "Röding Vollrads"	S84. Bonfol
37. Wintzenbach	139. Budenheim 1	S85. Moutier, Pâturage
38. Niederroedern, Hatten, Croetwiler	140. Nackenheim, Weisenau	S86. Court, Golat
39. Auenheim	141. Wiesbaden 1, Hessler, Kostheim	S87. Tramelan
40. Forstfeld	142. Flörsheim 1, Hochheim	S88. Tavannes
41. Oberseebach	143. Niederhöchstadt	S89. La-Chaux-De-Fonds, Le Locle
42. Büchelberg	144. Bommersheim	S90. Les Ponts de Martel
43. Région Bouxwiler-Grand Basberg	145. Beckersheim-Bodenheim	S90b. Les Verrières
44. Westhouse, Osthouse, Uttenheim	146. Niedererlenbach	S91. Balm
45. Sundhouse	147. Fechenheim	S92. Grenchen 1
45b. Ebertsheim	148. Rockenberg, Wölfersheim, Echszell	S93. Bennwil
46. Lahr, Dinglingen, Burgheim, Dingten,	149. Ostheim 1	S94. Vue Des Alpes
Mittersheim, Schutterlindenberg	150. Ravalzhausen (Mn2)	S95. Nebelberg
47. Nonnenweiler		
48. Mahlberg		
49. Erstein-Kraft	51. Pyrimont-Challonges, Génissiat	S100. Ornberg, Schwartz-Rüti, Güntisberg, Hüllstein
50. Allondans, Ste Suzanne, Pezol	52. Findreuse 3-27	S101. Esslingen
51. Bethoncourt	53. Chavannes	S102. Gerstel, Elgg, Schauenberg
52. Audincourt-Exincourt	54. La Chauda Fontaine, Rte de Serasson, Fornant 6-13	S103. Langnau
53. Rechésy	55. Lovagny	S104. Imenberg-Stettfurt, Chräzerentobel
54. Charnmoilles	56. Avanchet	S105. Rümikon
55. Bouxwiler-Oltingue-Kiffis	57. Choulex	S106. TMC-Glatbruggg
56. Wolschwiler	58. Vaux	S107. Hirschengrabentunnel Zh
57. Boecourt, Seprais	59. Veyron 2	S108. Andelfingen
58. Develier-Delémont	510. Cossonay	S109. Buchberg
59. Vicques	511. Boscéaz, Orbe	S110. Hellsigshausen-Fischbach, Ottenberg
60. Corban	511b. Eclepens, Alleveys, Mormont Entreroches	S111. Iberg
61. Laufen-Bryslach-Breitenbach	511c. Augine	S112. Küsnacht borehole
62. Oberwil-Therwil-Witterswil	512. Rances, Perrée, Valeyres	S113. Wiesholz
63. Leymen	513. La Chau	S114. Oeningen
64. Mettenberg	514. Sergy (SPM5)	S115. Greuterschberg, Mttelen-Weid
65. Knoeringue-Durlinsdorf	515. Ruisseau du Bey	S116. Schauenberg
66. Neuwiller, Allschwil	516. Talent 12	S117. Utzwil, Nutzenfluh
66b. Altkirch, Tagolsheim	517. Talent 1-19, Cuennet, Bavois	S118. Lauffenbach
66c. Dannemarie	518. Mt Chesau, Bois de Tey 5-7	
67. Sentheim	519. Region Lausanne: La Mèbre 630+698,	S120. Boswil 1
68. Schweighouse sur Thann	Bois-Genoud, Les Bèrgieres, Les Pierrettes,	S121. Hünenberg 1
69. Burnhaupt	Rochette, Paudèze 2-15, Flon 2, Le Lendar,	S122. Linden 1
70. Guewenheim, Mischelbach	Savigny, Chandelar B, Le Marcheret	S123. Fisch 2, Fischenbach section, Entlebuch 1
71. Cernay	520. La Cornalle	S124. Höhronen section
72. Hoelensheim-Lauw	521. Bellières	S125. Rigi section, Rigi 1 & 2
73. Froidefontaine, Bourogne,	522. Les Fontanettes 1-2	
Morvillars	523. La Roulavaz	
73b. Allenjoie	524. Broye 555	S130. Dessy
73c. Bethovillers & Roppe	525. Petite Chamberonne	S131. Molettes, Portettes, Lombardes, Palton
74. Reiningue	526. Vuibroye F. Moulin de Haut Crêt	S132. Covaye, Platté
75. Wittenheim, Sausheim	527. Cheseaux-Noréaz	S133. Quille, Pierredar, Lizerne, Diablerets
76. Staffelfelden, Wittelsheim	528. La Maugetta	S134. Tzanfeuron
77. Schoenensteinbach	529. Cheyres	S135. Fürstli, Schwand, Rothorn,
78. Rheinweiler	530. Le Gèrinnoz	Niesen, Sanetsch, Rawil
79. Zimmersheim, Rixheim,	530b. Vaulruz	S136. Foribach
Riedsheim	531. Arbogne 2	S137. Bürgenstock, Klimsenhorn
80. Brunnstadt, Hochstadt	532. Ameismühle, Gottéron 1+6	
81. Bruebach	533. Boudry	
82. Hartmannswiller, Wuenheim	534. Lieffrens	
83. Einsisheim, Ungersheim	535. Schiffenen 4	
84. Hornburg	536. Vully 1	
85. Habsheim	537. Brütelen 2	
86. Soultz Haut Rhin	538. Morigen	
87. Rouffach	539. Cortèbert	
88. Pfaffenheim, Hattstadt,	540. Sellgraben	
Herrlisheim, Gueberschwir	541. Messen	
89. Eguisheim	542. Verres, Corban	
90. Oberbergheim &	543. Basel-St Jakob	
Niederbergheim	544. Brochenen Fluh, Waldenburg-Humbel,	
91. Blodenheim, Fessenheim	Breitehöchi, Mümliswil	
92. Meyenheim	545. Oensingen	
93. Hirtzfelden	545b. Egerkingen	
94. Hettenschlag	546. Wolfwil, Wynau 1+2	

1. Steinbach/Gallensis Formation, Bürgen Fm., Klimsenhorn Fm.
2. Globigerinenmergel, Flysch sudhélvétique, Marnes à Foraminifères, Hohgant Formation (base)
3. Klimsenhorn Formation, Wildstrubel Fm.
4. Siderolithic, Eozäner Basiston
5. Calcaires à petites nummulites, Cerithium Beds, Couches à Cerithes, Couches des Diablerets, Sanetsch Fm.
6. UMM Meletta shale, Taveyannaz and Aldorf sandstones, Val d'Illiez Fm.
7. UMM Vaulruz Fm., Grigiser Mergel & Horwer Sandstein
8. UMM Top Baustein Beds
9. USM Molasse rouge, Untere Bunte Mergel and Conglomerates, Mt. Pelerin Fm., Beichlen Fm. and Uerscheli Fm., Homberg Beds, Losenegg Beds, Schwändibach conglomerate, Thun conglomerate, Honegg marls, Giltzschöpf conglomerate
10. USM Molasse à charbon and conglomerates
11. USM Calcaires and Dolomies
12. USM Grès and Marnes grises à gypse (GMGG)
13. USM Molasse grise de Lausanne (and local conglomerates), Gunten Quarzite, Granitische Molasse
14. OMM (Luzern Fm., St-Gall Fm., Sense Fm., Belpberg Fm.)
15. OSM (conglomerates)
16. Calcaires inférieurs, Krustenkalk
17. Obere Bunte Mergel (and sandstones)
18. OMM-OSM transition, Marnes rouges, Rote Mergel, Helicidenmergel, Graupensandrinne
19. OSM ("Oeningian") and Glimmersand
20. Hegau and Höwenegg series
21. Molasse alsacienne
22. Calcaires delemontiens
23. Calcaires de La Chaux
24. Karst Vue des Alpes
25. Gelberde, Raitsche
26. Fischeschiefer, Septarienton, Meeressand, Foraminiferenmergel,
27. Conglomerate de Porrentruy
28. Calcaires de Vermes
29. Anwiler Kalke
30. Juranagelfluh s.l., Vogesensande, Hipparionsande, Dinotheriensande, Bohnerzton, ?Wanderblock Fm.?, ?Heubergschotter?, Dorn-Dürkheim Beds, Lautersheim Formation
31. Sundgauschotter and Argiles de Bonfol, Weiße Serie
32. URG-conglomerates and breccias ("Steingang")
33. Planorbenkalke, Calcaires d'eau douce "Lutétiens", Bouxwiller Fm.
34. Lymneenmergel, Marnes vertes, Grüne Mergel
35. Melanienkalke
36. Salifère inférieure, Lower Salt Fm.
37. Messel Formation
38. Ältere Eisenberger Tonfolge, Eisenberger Grünton
39. Lower Pechelbronn Beds, Rote Leitschicht, Couches Rouge
40. Zone dolomitique
41. Salifère moyen, Middle Salt Fm.
42. Plattiger Steinmergel, Marnes et calcaires en plaquette, Zone fossilifère
43. Middle Pechelbronn Beds
44. Salifère supérieure, Upper Salt Fm.
45. Hausteine
46. Upper Pechelbronn Beds
47. "Serie grise", (Foraminiferenmergel, Fischeschiefer, Meletta-Beds, Cyrenenmergel, etc.), Alzey Fm., Bodenheim Fm., Stadelcken Fm., Sulzheim Fm. (pars)
48. Freshwater Beds, Niederroedern Beds
49. Tüllinger Kalk
50. Ältere Deckenschotter
51. Arvernensis-Schotter, Klebsand, Weißes Mio-Pliozän, Ältere Weisenauer Sande "Pliocene not differentiated", Pliozän I-III, Jungtertiär II (higher part)
52. Lower Cerithium Beds, Sulzheim Formation (pars)
53. Middle and Upper Cerithium Beds, Landschneckenkalke, Lower Hydrobia Beds, Oberrad Formation, Oppenheim Formation
54. Rüssingen Formation, Inflata-Beds and Wiesbaden Fm, Upper Hydrobia Beds, Landschneckenmergel, Niederrad Formation, Jungtertiär I
55. Praunheim Formation, Staden Fm., Bockenheim Formation, Salzhäusen Beds, Main Trapp, Jungtertiär II (lower part)

Fig. 6 List of lithostratigraphic units

Diem 1986; Kuhlemann and Kempf 2002; Lateltin 1988; Schlunegger et al. 1997). In the western part of the Swiss Molasse Basin (SMB), marine sedimentation stopped earlier (not younger than NP22, Berger 1992b; Berger et al. 1987; Carbonnel 1982) than in the eastern part, where marine sediments are known until NP24 or even NP25 (Doppler et al. 2000; Deutsche Stratigraphische Kommission 2002; Oberhauser et al. 1991). Following the general sea-level drop (corresponding with the Ch1 sequence at 28.5 Ma, see Hardenbol et al. 1998) these deposits are followed by the freshwater deposits of the Lower Freshwater Molasse (USM), which consist of conglomeratic alluvial fan deposits, fluvial sediments, and palustrine-lacustrine deposits, e.g., the *Molasse à Charbon* (see Berger 1992b, 1998; Engesser et al. 1984; Fasel 1986; Schlunegger et al. 1996, 1997a, b, c).

In the Plateau Molasse, some punctual lacustrine deposits of Rupelian age are known (*Calcaires inférieurs*), which are generally dated by charophytes (Berger 1992b; Weidmann 1984). In the Lower Chattian, fluvial

sediments (*Untere Bunte Mergel*) are present in the whole Foreland Basin. During the Middle and Late Chattian, lacustrine and brackish limestones, dolomites, and gypsiferous marls occurred. In the NE distal part of the basin, sedimentation probably started not before the middle Chattian (Müller et al. 2002; Schlunegger and Pfiffner 2001). All mentioned units are relatively well dated by mammals, charophytes, otoliths and magnetostratigraphy (Berger 1992b, 1996; Engesser 1990; Engesser and Mödden 1997; Kempf et al. 1999; Kissling 1974; Platt 1992; Reggiani 1989; Reichenbacher and Weidmann 1992; Reichenbacher 1996; Schlunegger et al. 1996, 1997a, b, c; Strunck and Matter 2002).

Miocene and Pliocene

No Miocene sediments have been recorded in the subalpine Molasse of Western Switzerland. Some conglomeratic fans persist in Eastern Switzerland until the Middle Aquitanian (Schlunegger et al. 1997a, b, c).

In the Plateau Molasse, fluvial facies characterizes the Aquitanian deposits (*Molasse Grise de Lausanne*, *Obere Bunte Mergel*, *Granitische Molasse*, see Berger 1985; Keller et al. 1990). A marine transgression (partly corresponding to the global sea-level rise following the Bur 1 sequence at 20.5) flooded the Foreland Basin during the late Aquitanian (MN2) and Burdigalian; its typical deposits are the Upper Marine Molasse (OMM) (Berger 1985, 1996; Graf 1991; Homewood and Allen 1981; Keller 1989; Kempf et al. 1997, 1999; Kempf and Matter 1999; Kuhlemann and Kempf 2002; Schaad et al. 1992; Schoepfer 1989; Strunck and Matter 2002).

During the Langhian, marine sediments were confined to the North, with local "Juranagelfluh" in the NE (Graf 1991); whereas fluvial sediments related to alluvial fans were present in the South (OSM = Upper Freshwater Molasse), and persisted until the Serravallian (MN7, ?MN8) (Bolliger 1992, 1997, 1998; Kälin and Kempf 2002; Kempf et al. 1999).

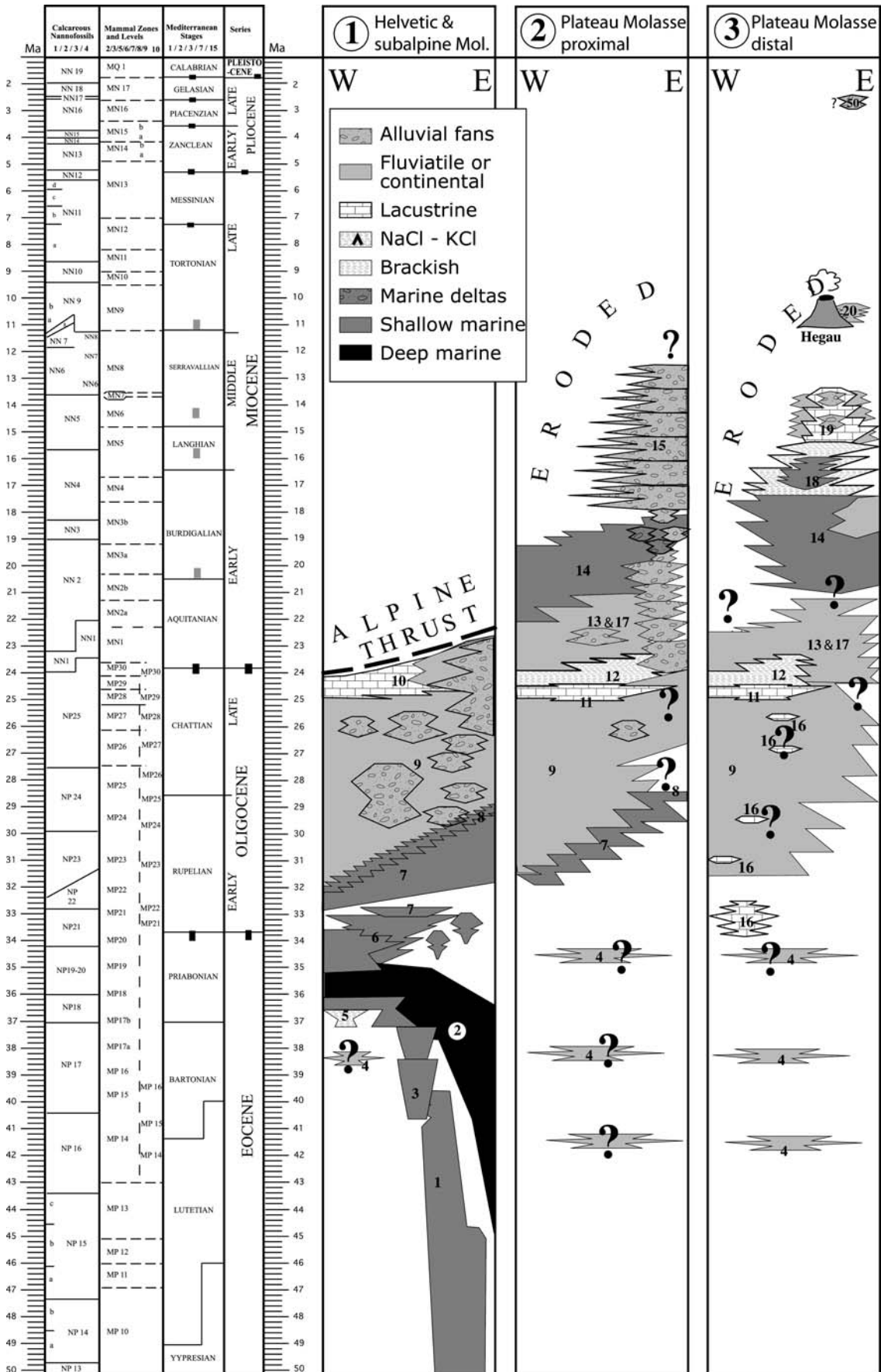
No Late Serravallian, Tortonian and Messinian sediments have been recorded, except in the Northeastern part (Höwenegg, MN9, Tobien 1986). However, several authors suggested an additional 700 m and up to more than 2,000 m were deposited in the Eastern and Western Molasse basin, respectively (see Kuhlemann and Kempf 2002).

The Pliocene is represented only by conglomerates ("Ältere Deckenschotter", Graf 1993) in the Northeastern part, with one dated locality (Irchels, MN17, see Bolliger et al. 1996).

The Molasse of the Jura Mountains (Fig. 8)

In the Jura mountains, Tertiary sediments have been preserved in a number of synclines.

These series were studied in detail by Picot (2002) and Becker (2003).



◀
Fig. 7 Stratigraphy of the Helvetic domain, Subalpine and Plateau Molasse

Eocene

Eocene deposits are represented by the *Siderolithic*. Even though it was never precisely dated, its Lutetian, Bartonian and Priabonian age is suggested owing to correlation with other siderolithic localities (see Chap 2.1)

Oligocene

During the Oligocene, the different synclines experienced a diversified sedimentary history.

In the Southwest (Fig. 8, column 4), some lacustrine limestones (*Calcaires inférieurs*) appear, which, so far, have not yet been precisely dated. These are overlain by fluviatile sediments, probably of Rupelian to Chattian age (*Untere Bunte Mergel*), and finally by calcareous limestones and gypsiferous marls of Late Chattian to Early Aquitanian age (Berger 1996; Mojon et al. 1985).

In the Northwest (Fig. 8 column 5), no deposits are known from that time.

In the central part, the conditions vary from South to North:

- in the South (Fig. 8 column 6), lacustrine limestones are present that range in age from Early Rupelian (dated by charophytes of the *tuberculata* zone in the Moutier syncline, see Reichenbacher et al. 1996) to Early Chattian (mammals in Oensingen, Wynau or Soultce, see Engesser and Mödden 1997). They are overlain by the fluviatile *Molasse alsacienne* with a possible marine horizon, but a reworking of marine alpine material cannot be excluded (Reichenbacher et al. 1996). In the Late Chattian, lacustrine limestones (*Calcaires delémontiens*) and gypsiferous marls were common and persisted until the Early Aquitanian.
- in the North (Fig. 8 column 7), Tertiary started in the Priabonian with lacustrine-brackish deposits (*Gelberde*, *Raitsche*) in the Delsberg Basin, and with the *Porrentruy conglomerates* near the French border (see also Clement and Berger 1999; Picot et al. 2004). These deposits are overlain by marine Late Rupelian sediments (*Foraminiferenmergel*, *Fischschiefer*, *Septarienton*), which additionally build up the base of the Tertiary deposits in the Laufen Basin (NP22 and 23). Then follow fluviatile sediments (*Molasse alsacienne*), which are dated as MP23 to MP26, depending on the presence or absence of the marine sediments (Clement and Berger 1999). During the Late Chattian and Early Aquitanian, the *Calcaires delémontiens* indicated lacustrine sedimentation (Picot et al. 1999).

In the eastern Jura (Fig. 8 column 8), we find again the succession of the *Calcaires inférieurs*, the *Molasse*

alsacienne, and the *Calcaires delémontiens*. The latter are well dated (from MP29 to MN1) in the Brochene Fluh section (Engesser and Mayo 1987; Becker et al. 2001).

Miocene and Pliocene

In the Southwest (Fig. 8 column 4), the Aquitanian is only represented by the uppermost part of gypsiferous marls (dated as MN1), and the *Calcaires de La Chaux*, which is a famous fossil locality belonging to MN2 (Engesser and Mödden 1997). The following discordant transgression of the OMM is probably of Burdigalian age. Younger sediments are not known in this area.

In the Northwest (Fig. 8 column 5), Tertiary sedimentation started with the OMM. The overlying brackish-marine *Marnes rouges* are dated as NN4-5. Then follows the lacustrine *Oeningian* facies (MN6-7, see Kälin et al. 2001).

In the central part (Fig. 8 column 6 and 7), there is a time gap between the Early Aquitanian (MN1, top of *Calcaires delémontiens*) and the overlying transgressive OMM, which in the Tavannes area belongs to the Burdigalian (MN3, De Beaumont et al. 1984) and in Glovelier to the late Burdigalian (MN4, Hug et al. 1997). Thus, the Middle and Late Aquitanian, and perhaps the earliest Burdigalian sediments are missing. The OMM is overlain by the lacustrine *Oeningian* facies (MN6-7) or the *Calcaires de Vermes* (MN5, Kälin 1993; Kälin and Kempf 2002). Then follow conglomerates and sands (*Jura Nagelfluh*, *Hipparion-* and *Vogesensande*, Bois de Raube Formation), which have been dated by mammals to MN6 up to MN9 (Kälin 1993, 1997b; Kälin and Engesser 2001). The so-called *Wanderblock* Formation probably was deposited during this time (Kemna and Becker-Haumann 2003).

In the Eastern part (Fig. 8 column 8), the OMM is generally followed by the brackish complex of the *Marnes rouges*, and then by the fluviatile “Glimmersandschüttung”. During the Burdigalian, the “Graupensandrinne”, an estuarine system running along the distal border of the German part of the Foreland Basin, was active (Reichenbacher et al. 1998).

The only Late Miocene or Pliocene deposits in this area are the karstic filling of the Vue des Alpes (which is dated as MN15 and is posterior to the main Jura folding, see Bolliger et al. 1993), and the Höhere Deckenschotter from Irchels (Bolliger et al. 1996).

The South and Middle Upper Rhine Graben (URG) (Fig. 9)

Eocene

Eocene sedimentation generally began with the rarely dated Siderolithic (Pharisat 1982, MP19) that can be attributed to the Middle to Late Eocene in comparison

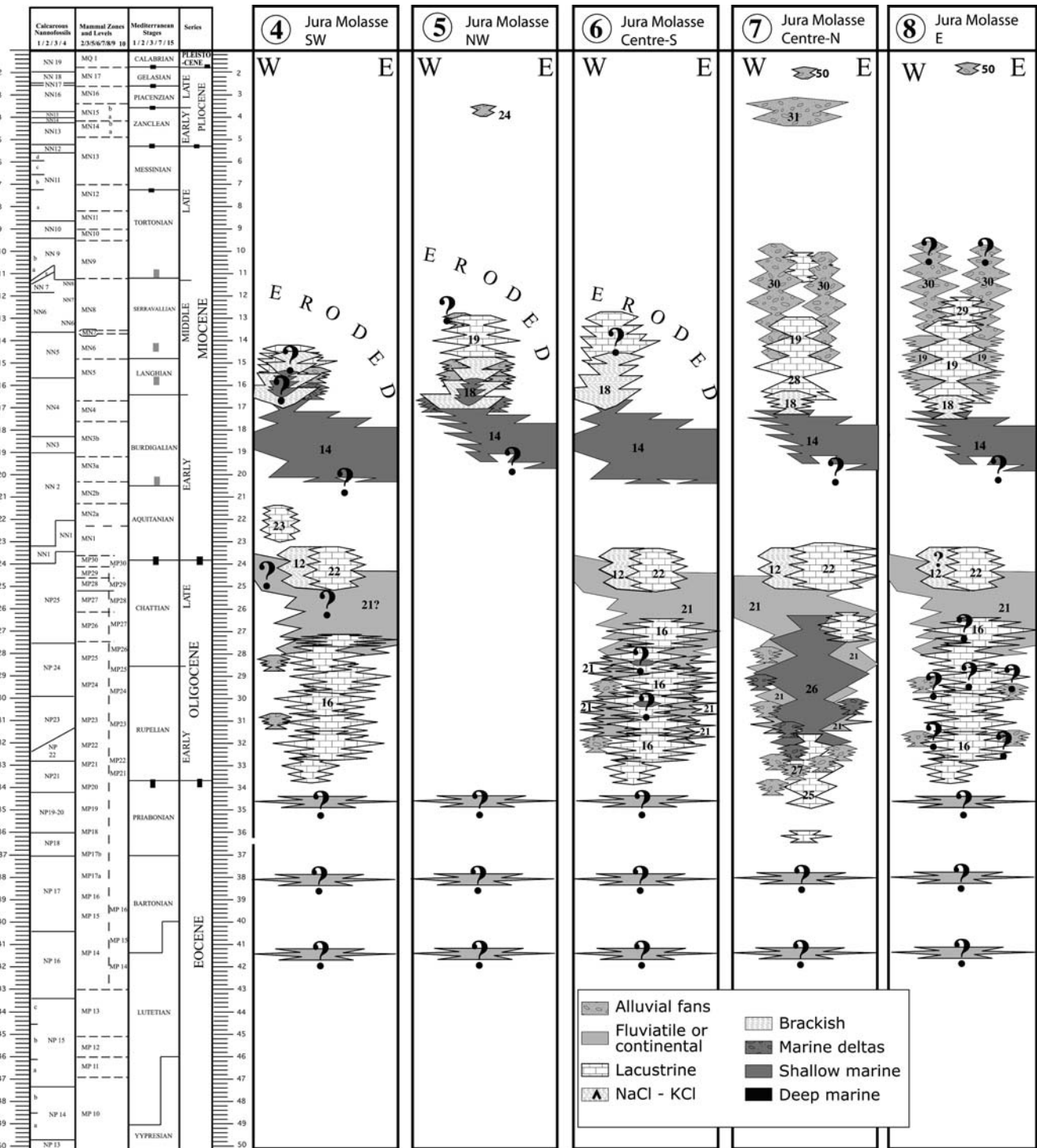


Fig. 8 Stratigraphy of the Jura Molasse

with adjacent areas. It is followed by various sedimentary facies depending on the syntectonic processes in the URG at that time.

In its Southern part (Fig. 9 column 9), the URG is tectonically subdivided W to E into the following four major areas, the Dannemarrie Basin, Mulhouse Horst (with Altkirch- and Sierentz blocs), Sierentz Basin, and the Basel Horst (Förster 1905; van Werweke 1908;

Wittmann 1949; Doebel 1970; Lutz and Cleintuar 1999; Sittler 1965, 1972, 1992; Sittler and Schuler 1988):

- In the Dannemarrie Basin, important accumulation of brackish sediments (*Lymnaeaenmergel*) and salt (Lower Salt Formation) is known. Their ages span probably from the Lutetian to the Priabonian, possibly from NP12 to NP17 (see Chateaufneuf 1983; Sittler 1965;

Schuler 1983, 1988, 1990). However, it is not clear if sedimentation was continuous, or a gap corresponds to the Bartonian, as postulated by Sissingh (1998). In the western part, some undated conglomerates are tentatively assigned to the Lutetian and the Priabonian (Düringer 1988).

- On the Mulhouse Horst, brackish marls and lacustrine limestones are present, probably ranging from Lutetian to Priabonian. The *Melanienkalke* are clearly dated to the Priabonian (MP18) at Brunstatt (Schwarz 1997; Tobien 1949, 1968, 1987, 1988a).
- In the Sierentz Basin, the *Lymnaea marls* (also called *Green marls*), lacustrine limestones and the Lower Salt Formation (but without salt!) were deposited (Sittler 1965). In the eastern part, some conglomerates probably accumulated during the Priabonian. At the locality Pfaffenweiler, they could be dated by mammals to MP18 (Tobien 1968, 1987, 1988b).
- In the northern part of the Basel horst, a condensed sequence is present, which is very similar to that of the Mulhouse horst. In the southern part, Eocene sediments are absent (Doebel 1970).

In the middle part (Fig. 9 column 10) we recognize a similar situation from W to E, with lacustrine limestones, *Lymnaea marls* and Salt formation (especially developed in the Western part). A brackish complex (*Zone dolomitique*) generally overlies these deposits, but is not dated. Conglomeratic facies is present at both borders (particularly from the Priabonian onwards).

Toward the north, the Lower Pechelbronn Beds are partly dated by mammals as Priabonian (MP 20, Tobien 1987, 1988).

In the northern part (Fig. 9 column 11) a Lutetian age of the lacustrine limestones at Bouxwiller is documented by charophytes and mammals attributed to the MP13b zone (Jaeger 1971; Schmidt-Kittler et al. 1997; Biochrom'97 1997). Also the *Lymnaea marls* (or *Zone de marnes à anhydrite*) could be correlated by charophytes with the Lutetian (*embergeri* zone, Breuer and Feist 1986; Riveline 1985). An undated brackish complex (*Zone de marnes dolomitique*) generally overlies these deposits. The *Zone de marnes à anhydrite* and the *Zone de marnes dolomitique* are combined by Schnaebel (1948) to the *Zone dolomitique* (or *Green Marls*), which underlie the Lower Pechelbronn Beds in the Pechelbronn Basin). Conglomerates are also present on both sides of the graben.

Oligocene

In the southern part (Fig. 9 column 9), the Early Rupelian sedimentation is represented by

- the Middle and Upper Salt Formation (Dannemarie Basin)
- the *Zone fossilifère* and the *Streifige Mergel* (= *Marnes en plaquettes*) as well as the *Haustein*. The

latter is dated as NP21-22 (Schuler 1988, 1990) and MP21 (Altkirch, see Storni 2002, Mulhouse Horst) – the conglomeratic facies at both borders.

The *Zone fossilifère* represents the first Rupelian transgression from the North Sea (partly corresponding with the global sea-level rise following the Ru 1 sequence at 33 Ma, see Haredbol et al. 1998) and separates the Middle from the Upper Salt Formation.

In the Middle Rupelian, the most important marine transgression, second marine Rupelian transgression from the North Sea, corresponding to the global sea-level rise between sequences Ru2 and Ru3, see Haredbol et al. 1998) flooded the whole basin, with the classical succession: *Foraminiferenmergel*—*Fischschiefer*—*Meletta*—*Schichten* (partly). These sediments range from NP23 to NP24 (Grimm 2002a; Müller 1988; Pharissat 1991). They are combined with the Late Rupelian brackish-marine *Cyrenenmergel* to the so-called *Graue Serie* (= *Série grise*) and represent an important marker on seismic lines.

During the Early Chattian, the fluvial sediments of the *Molasse alsacienne* interfingered to the north (Tüllinger Berg) with lacustrine or brackish marls (*Gipsmergel*, see Wurz 1912). During the Middle and Late Chattian, the lacustrine limestones of the *Calcaires delémontiens* and the *Tüllinger Kalk* were deposited. The *Calcaires delémontiens* are well dated as MP29 to MN1 in the Jura Molasse (Becker 2003; Picot 2002), and the *Tüllinger Kalk* can be correlated by charophytes to the Middle Chattian *Stephanochara ungeri* Zone (Nötzold 1962).

In the middle part, the Middle Salt Formation and the brackish-marine Middle Pechelbronn Beds are present. Even though their correlation is not fully solved, they probably belong to the NP22 and MP21 zones (Martini 1973; Griessemmer 1998, 2002; Storni 2002; Derer 2003 and discussion on Fig. 2).

The overlying Upper Salt Formation and the Upper Pechelbronn Beds can be assigned to MP21 and MP22 (= ? NP22 and NP23, Schwarz and Griessemmer 1992).

The marine-brackish Middle to Late Rupelian *Série grise* covers these units, and is itself overlain by the fluvio-lacustrine Niederroedern Beds and the brackish-lacustrine Cerithium Beds (Doebel and Geissert 1971; Doebel et al. 1976). In some localities, these freshwater sediments are also referred to as *Molasse alsacienne*.

In the northern part (Fig. 9 column 10 and 11), the Early Rupelian sedimentation is represented by the Middle Pechelbronn Beds, followed by the Upper Pechelbronn Beds, the *Série grise* and the Niederroedern Beds. Based on lithostratigraphic correlations with the Süßwasserschichten of the Mainz Basin, the Niederroedern Beds belong to the Middle to Upper Chattian. The overlying Lower and Middle Cerithium Beds, and the lowermost part of the Upper Cerithium Beds are of Upper Chattian age (Martini 1978; Reichenbacher 2000).

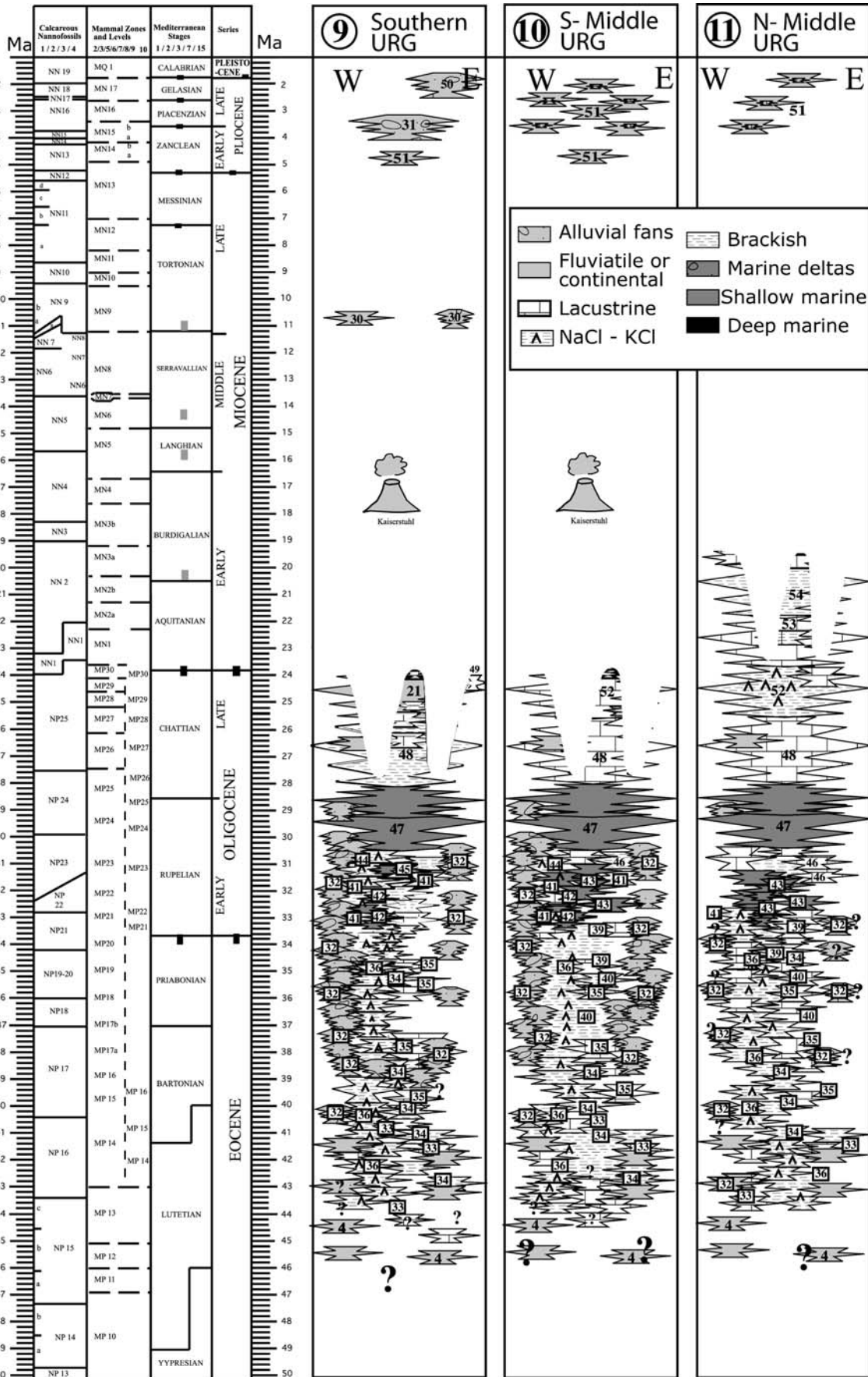




Fig. 9 Stratigraphy of the South and Middle URG

Miocene and Pliocene

From the Aquitanian onwards, the southern and middle parts of the URG were subjected to erosion. Sediments may have been deposited until uplift of the Vosges and Schwarzwald (Burdigalian), or the nondeposition already commenced during the Middle and Late Aquitanian, as proposed for the Jura molasse (see Picot 2002; Becker 2003).

Sedimentation resumed only locally during the Early Tortonian, with the deposition of the *Dinotherium* and *Hipparion sands* (dated by mammals as MN9, Tobien 1986a, 1988b).

No Messinian sediments are known. The last sedimentary event is represented by the *Sundgau Schotter*, which could be attributed to the Middle Pliocene (MN15 to 16) in comparison with the Bresse Graben mammal localities (Petit et al. 1996). The *arvernensis Schotter* is also punctually represented and dated as MN14–MN16 (Tobien 1988b). Pliocene deposits (dated by spores and pollens) are also known in the vicinity of Colmar and Strasburg.

In the northern part (Fig. 9 column 11), sedimentation continues in the Aquitanian with the brackish Upper Cerithium Beds, which are overlain by the Inflata Beds (now Rüssingen Formation, see chapter Miocene and Pliocene), which are followed by the Lower Hydrobia Beds (now Wiesbaden-Formation, see chapter Miocene and Pliocene). This sequence seems to belong to the Aquitanian, according to its correlation with the Mainz Basin (MN1 and MN2, see Reichenbacher 2000). The Upper Hydrobia Beds are still present, and may be attributed to the Early Burdigalian (MN3) by correlation with the Hanau Basin (see chapter Miocene and Pliocene). After an important gap (Burdigalian to Messinian), the sedimentation resumed again during the Pliocene, as known in the Hagenau area.

The northern URG, the Mainz Basin and the Hanau Basin (Fig. 10)

Eocene

During the Eocene and Early Oligocene the northern part of the URG was not yet developed. Instead of this the main structures were influenced by preTertiary tectonism. From north to south the following perift structures are recognized, namely the Rüsselsheim Basin, Palatinat-Stockstadt Ridge and Marnheim Bay (Grimm and Grimm 2003).

In the Rüsselsheim Basin, the Marnheim Bay and the southward connected northern URG sedimentation began with the terrestrial *Eozäner Basiston* and *Basissand* (Sonne 1968; Schäfer 1996, 2000). On the Palatinat-Stockstadt Ridge, especially on its eastern part (the

so-called Sprendlingen Horst), the Messel Formation accumulated in a maar-lake (Harms et al. 2000, 2001; Jacoby et al. 2000). These sediments are dated by mammals as MP11 (Tobien 1988b).

The Priabonian Lower Pechelbronn Beds overlay Eocene basal sediments in the Rüsselsheim Basin. In the Marnheim Bay and in the URG, these basal sediments are overlain by the lacustrine to brackish deposits of the *Green marls*, the *Rote Leitschicht*, and the *Eisenberg clays* (Schäfer 2000; Grimm and Grimm 2003). The latter can be correlated to the Middle Lutetian by pollens and spores (Hottenrott 2000, 2002). The *Green marls* correlate to the Lutetian *embergeri* zone and the Priabonian *vectensis* zone by charophytes (Schwarz 1997). Generally, the Lower Pechelbronn Beds are dated as MP 20 (Derer 2003; Gad et al. 1990; Gaupp and Nickel 2001; Schwarz 1993, 1997; Schwarz and Griesemer 1994, Tobien 1949, 1968, 1987).

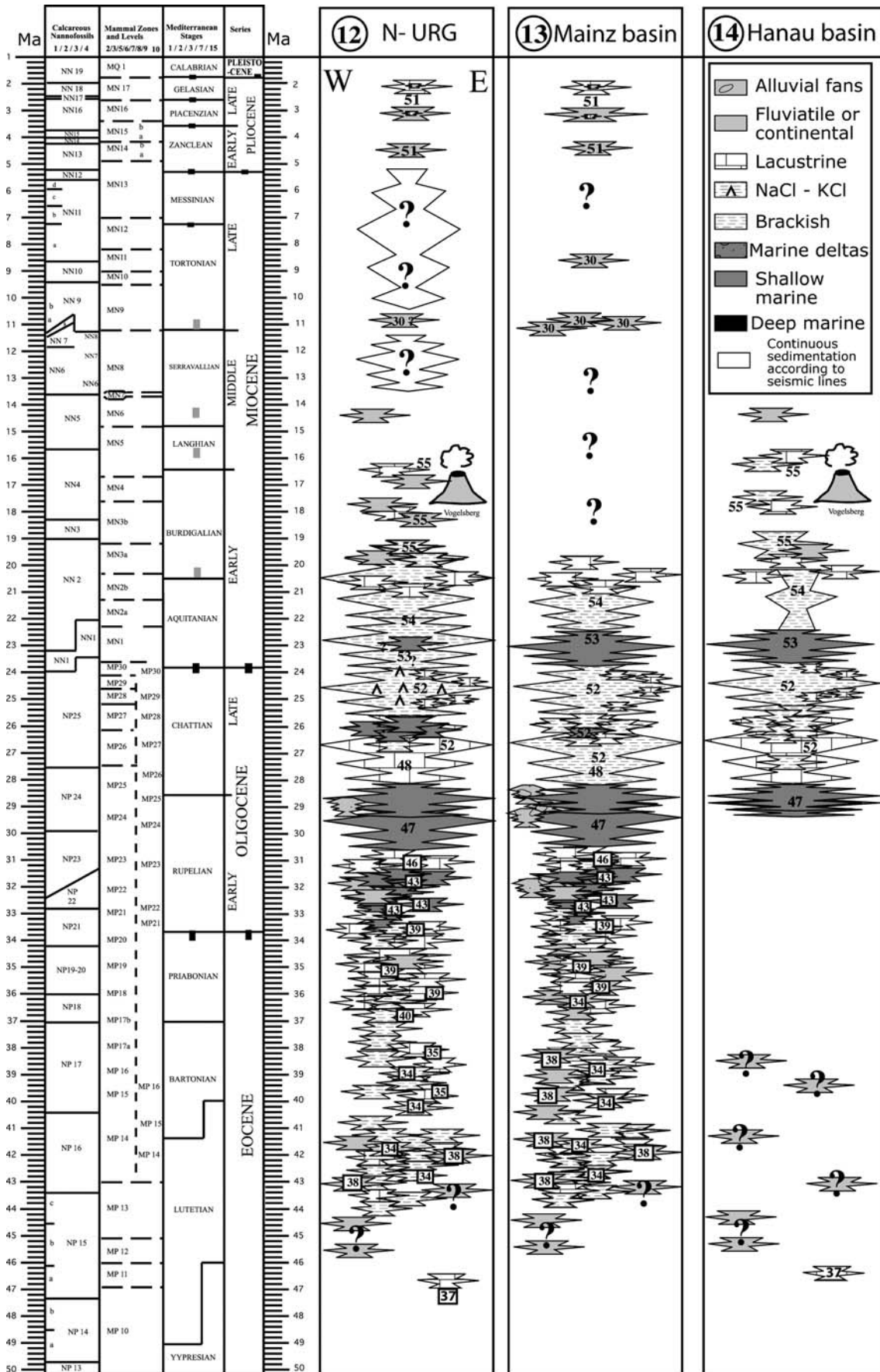
Along the border of these basins, terrestrial alluvial fan deposits of the “Steingang” accumulated during the Lutetian to Early Rupelian (Grimm and Grimm 2003). Eocene sediments are absent in the Hanau Basin.

Oligocene

The Middle Pechelbronn Beds have been correlated with the first brackish-marine Rupelian transgression (Sissingh 2003). The Palatinat-Stockstadt Ridge was partly flooded, and the Rüsselsheim Basin was connected to the URG (Grimm and Grimm 2003). The Middle Pechelbronn Beds are well dated by nannoplankton (NP 22 according to Martini 1973, 1982, 1990; Martini and Reichenbacher 2004) and foraminifers (NSR 7a according to Grimm 2002b; Grimm et al. 2005). In the central parts of these basins, the Middle Pechelbronn Beds are overlain by the fluvial Upper Pechelbronn Beds. The Pechelbronn-Group was not deposited in the Hanau Basin (Hottenrott 1988).

The second and third Rupelian marine ingressions, which invaded the whole area (corresponding to sea-level highstands between Ru2/R3, and Ru3/Ru4 respectively) are dated to the top of NP23 up to NP24 (Doebel and Sonne 1974; Doebel et al. 1980; Grimm 1998, 2002a; Grimm et al. 2002; Grimm and Grimm 2003; Hottenrott 1988; Martini 1982; Müller 1988; Pross 1997; Tobien 1987). Several formations have been recently introduced for these sediments (Alzey-, Bodenheim-, Stadelcken-, Sulzheim Formation, Grimm et al. 2000). At this time the URG transected the older structures and reconfigured the investigated area into the western Mainz Basin, the central URG and the eastern Hanau Basin. The Hanau Basin was first flooded during the deposition of the upper part of the Bodenheim Formation (= Oberer Rupelton).

Following regression of the sea (corresponding to the sea-level drop at the sequence Ch1 and/or Ch2, see Hardenbol et al. 1998), brackish to fluviolacustrine sediments dominated the area (Niederroedern Schichten,



◀ **Fig. 10** Stratigraphy of the North URG, Mainz- and Hanau basins

Süßwasserschichten = Sulzheim Formation pars), dated to MP24 (Mödden et al. 2000; Reichenbacher 2000). These are overlain by (or pass laterally to) the brackish Lower Cerithium Beds, which have been correlated to the second Chattian transgression (Schäfer 1996, Sissingh 2003). At the same time alluvial fan deposits of the Budenheim Formation accumulated along the southwestern border of the Rhenish Slate Mountains (Schäfer and Kadolsky 1998). The Lower Cerithium Beds are overlain by (or pass laterally to) the brackish-marine sequence of the *Landschneckenkalk*, the Middle Cerithium Beds and the lowermost part of the Upper Cerithium Beds. This sequence still belongs to the Late Chattian (Martini 1978; Kadolsky 1988; Reichenbacher 2000; Schäfer 1988).

Miocene and Pliocene

During the sedimentation of carbonate platform and lagoonal sediments in the border areas of the URG (= Mainz Group according to Grimm and Grimm 2003) a basin facies was formed in the URG (e.g., Rothausen and Sonne 1984; Sissingh 2003).

The Upper Cerithium Beds span the Chattian–Aquitania boundary. The brackish-lagoonal sediments of the Oberrad Formation (= upper part of the Upper Cerithium Beds, see Schäfer and Kadolsky 2002) correlate, owing to mammals and otoliths, with the Aquitanian (MN1, MN2; Engesser et al. 1993; Mödden 1996; Schäfer and Kadolsky 2002; Försterling et al. 2002; Försterling and Reichenbacher 2002).

The overlying brackish-lacustrine marls and limestones of the Rüssingen Formation (= *Inflata*-Beds) still belong to the Aquitanian (MN2a, Engesser et al. 1993). In the uppermost part of the Rüssingen Formation, biota indicates a further brackish-marine incursion from the south (Schäfer 1984; Reichenbacher 2000; Sissingh 2003), which is again followed by an incursion from the North Sea (Martini 1981, 2000). This incursion (probably corresponding with the major sea-level highstand following the Aq 1 sequence) characterizes the base of the Wiesbaden-Formation (= Lower Hydrobia Beds, see Reichenbacher and Keller 2002), which consists of mainly bituminous marls and limestones and is of Aquitanian age, except perhaps its uppermost part.

During the Burdigalian, brackish sedimentation was progressively reduced, and replaced by lacustrine and fluvial deposits. In the Hanau Basin, the Upper Hydrobia Beds, the Niederrad Formation (= *Landschneckenmergel*), and the Praunheim-Formation (= *Prosothenia*-Beds) were deposited (Grimm and Hottenrott 2002; Hottenrott 1988; Radtke and Kümmerle 2005; Reichenbacher 2000). In the URG the “Jungtertiär I” was formed. The Niederrad Formation was correlated to MN3 by Stephan-Hartl (1972) and Tobien (1987).

During the Langhian and the Early Serravallian, sedimentation was essentially continental, brackish or lacustrine in the northern URG and the Hanau Basin (Staden Formation, Bockenheimer Fm., lower part of the “Jung Tertiär II”, see Grimm and Hottenrott 2002; Hottenrott 1988). A plateau-basalt layer (“Maintrapp”), which has a radiometric age of 16.3 Ma, interfingers between the Staden Formation and Bockenheimer Formation (Fuhrmann and Lippolt 1987; Grimm and Hottenrott 2005). There is no sedimentation known in the Mainz Basin from Late Burdigalian to Late Serravallian. Dated Middle and Late Serravallian sediments are unknown in the whole area.

The presence of limno-fluvial Tortonian deposits is attested by the Lautersheim Formation, the *Dinotherium sands* and the Dorn-Dürkheim Formation, which are dated at several localities of the Mainz Basin to MN9 and MN11 (Franzen 1997, 2000; Franzen and Storch 1975, Grimm and Grimm 2003; Lutz et al. 2003; Rothausen and Sonne 1984; Schäfer 2000; Tobien 1980, 1988b, 1986b).

Uppermost Miocene to Pliocene sediments are known from some areas in the Mainz Basin and the western part of the northern URG and are represented by the “White Mio-Pliocene” and Bohnerz-clays (Rothausen and Sonne 1984; Grimm and Grimm 2003).

Piacenzian fluvial sediments are represented by the *Arvernensis-gravels* and *Weisenau sands* of the northernmost URG and the Mainz Basin and dated by magnetostratigraphy and heavy mineral associations (Fromm 1986; Semmel 1983). The late Pliocene probably exists at least in the Heidelberg and Frankfurt area (Hottenrott et al. 1995).

The presence of several unconformities (between early Serravallian and Tortonian, between Tortonian and Lower Messinian and below the Piacenzian *arvernensis-gravels*) is still subject to controversy: field and borehole observations indicate three breaks in sedimentation (this study), whereas the new river seismic lines do not show any unconformity during the Mio-Pliocene in the northern part of the URG (Dèzes et al. 2004).

Conclusion

The present study gives a general stratigraphic frame for the Swiss Molasse and the URG deposits. However, several points are still hypothetical and have to be clarified in the future:

- precise dating of Eocene to Early Oligocene deposits is still unclear, particularly concerning the Salt Formation, the *Green marls* and the Pechelbronn Group; better data probably will be accessible in the near future due to the study of borehole material from the Pechelbronn and Colmar areas, which are still available thanks to the BRGM.
- the stratigraphic relationships between the Middle Pechelbronn Beds, the *Streifige Mergel*—*Hau-*

steincomplex, the Jura Molasse and the eastern distal Molasse Basin have to be clarified.

- the problems of unconformities (Bartonian, Burdigalian in the southern part, Middle to Late Serravallian and Lower Messinian for the whole basin) have to be checked by comparing new (and old) seismic lines.
- the precise dating of the Pliocene sediments is still unclear, particularly in the Middle and Northwestern URG.

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